

# TYPHOONS DRIVEN MORPHODYNAMICS OF THE WAN-TZU-LIAO SAND BARRIER (TAIWAN).

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This study focuses on the dynamics of the Wan-Tzu-Liao sand barrier when it is struck by one typhoon, a group of typhoons or the whole season of typhoons and monsoon. In the framework of the KUN-SHEN project, 7 months of monitoring (2011-2012) provided 20 topobathymetric surveys (within a 200 m long segment of the barrier from the subtidal zone to the back-barrier) and acquisitions of offshore, nearshore and shallow water hydrodynamics including velocity profiling, free surface measurement and absolute pressure. Offshore waves were extracted at Cigu buoy (18 m of water depth). Nearshore waves were acquired from the current profiler deployed 400 m off the coast in 4 m of water depth and water level on the subaerial beach were acquired from two pressure sensors deployed in the lower part of the subtidal zone and the dune crest. Morphologic changes of the emerged beach were monitored using D-GPS each week during winter monsoon season and just before and after each event during summer typhoons season. Amongst the eight studied typhoons, TALIM is the most energetic event observed. Offshore wave height reached  $H_s = 10.34$  m ( $T_p = 14.6$  s) at Cigu buoy and  $H_s = 2.3$  m ( $T_p = 13.4$  s) at the profiler. Surface waves dynamics throughout surf and swash zones are progressively dominated by infragravity motions. A part of the dune is surged at the storm apex. Morphological changes include 6.7 m of dunefoot retreat and a sand transfer from a dune breach to wash-over deposits in the lagoon. More surprisingly, the foreshore was nourished ( $2261 \text{ m}^3 \pm 268 \text{ m}^3$ ) as well as the whole sand barrier ( $+1920 \text{ m}^3 \pm 1071 \text{ m}^3$ ). The sand input is concomitant to the redistribution of sand from the intertidal sandbars interpreted as the ultimate stage of the subtidal bars landward migration. Although winter are erosive season ( $-4995 \text{ m}^3 \pm 1071 \text{ m}^3$ ), the summer results in an accretion period ( $3556 \text{ m}^3 \pm 1071 \text{ m}^3$ ) with a shoreline seaward shift about 10.4 m. Over the annual time period, the sand barrier recorded 18.4 m of retreat coupled with a small sand loss ( $-1439 \text{ m}^3 \pm 1071 \text{ m}^3$ ) without any significant abrasion of the dune-top.

*Keywords: typhoon; moderate/extreme storm-impact; morphodynamics*

## INTRODUCTION

Beach morphodynamics have received an increasing interest during past decades, see e.g. (Komar, 1998; Short, 1999; Dean and Walton, 2006; Dronkers, 2005). Over a wide spectrum of time scales, the sedimentary beach dynamic is controlled by several factors including meteorologic and hydrodynamic forcings (wind, atmospheric pressure, wave, instantaneous sea level and tide), sand size/composition, initial morphologies, geological settings and mean sea-level history (Coco et al., 2014). Ruggiero et al. (2005) highlighted the importance of improving our understanding of beach behaviour associated with storms on several time-scales, including the interaction between environmental forcing and coastal morphology occurring over a wide range of time (from second to decades) and space scales (from centimeter to kilometers). However, interactions and feedback between the maximum excursion of waves on the beach (run-up) and emerged beach changes have been shown to play a great role in the beach response to storms. Four conceptual model regimes have been proposed by Sallenger (2000). The *swash regime* driven by mildest conditions results in a moderate setup at the shore and morphologic changes of the berm. In this case, the sand eroded is generally moved offshore. The *collision regime* occurs during more severe wave conditions and induces foreshore and dune-front erosion by scarping. The dune surge process displays the *overwash regime* resulting from severe elevation of the sea level under heavy energetic conditions induced a landward sand transport. *Inundation* is the final and most extreme regime and occurs when the beach and dunes are completely and continually under water, leading to dune erosion and/or island breaching. Based on these models, several studies helped to detail the relationships between sea level and morphologic responses of the beach when the system is forced by paroxysmal wind and waves conditions (hurricane in U.S coasts or typhoon in the Western Pacific) (Stone et al., 1993; Morton et al., 1995; Stone et al., 2004; Donnelly et al., 2004; Houser et al., 2008; Qi et al., 2010; Stockdon et al., 2007; Plant and Stockdon, 2012; Masselink and Heteren, 2013; Roberts et al., 2013).

In this context, our study aims at investigating the interactions between wave, water level and morphology of a sand barrier driven by paroxysmal conditions over time scales ranging from event to seasons. The main part of the paper presents field hydro-morphodynamical data from a single typhoon to the entire

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typhoon and monsoon seasons. In the discussion, we propose a new conceptual model of beach morphodynamics subjected to monsoon and typhoons conditions.

## SETTINGS

### Geodynamics context

In Taiwan, the seismic activity, the rapid rates of convergence, the deformation and erosion reveal a highly active tectonic setting, driving severe geological hazards such as landslides or earthquakes (Suppe, 1984; Lallemand et al., 2013).

Regarding morphology of the island, strong discrepancies occur between East facing the Pacific Ocean and West along the Taiwan Strait. The East coasts consists of rocks and gravel and high mountains (many of them are more than 3000 m high) plunging down into the sea with a steep slope (slope in order of 1 in 10; (Ou et al., 2002)), resulting in 4000 m of water depth less than 50 km off the coast. Conversely, the western coastal plains, which are our areas of interest, consist of finer sand with a very mild slope (slope in order of 1 in 200 to 500; (Ou et al., 2002)) due to the occurrence of the Taiwan strait. Taiwan straight is 350-400 km long, 180-200 km large and less than 60 m deep continental shelf between Taiwan and mainland China (Lin et al., 2005; Liao et al., 2008; Zhang et al., 2013).

### Hydro-morphodynamics context

Taiwan Island is exposed both to geological tectonic hazards (landslides and earthquakes) and extreme meteorological events. The wind and wave climates are governed by the alternating monsoon and typhoons seasons (Jan et al., 2002). We focus here on the hydro-morphodynamic context of the south-western coast of Taiwan. The monsoon season (from October to March) drives a roughly constant North wind of 5-6 m/s inducing southward alongshore drift. Significant wave height ( $H_s$ ) ranges from 1 to 1.5 m and wave period ranges between 6 and 8 s (Liao et al., 2008; Chen and Kuo, 2008). Summer typhoons season (from May to September) is a mix between long fair weather periods and short but highly energetic conditions during typhoons and tropical storms. The former are associated to small waves ( $H_s < 1$  m,  $4 < T_p < 7$ s) (Zhang et al., 2013) while the six to eight typhoons crossing or passing close to the island each year (Lin, 2013; Zhang et al., 2013) drive extreme winds, waves (32 m during the typhoon Krosa in 2007 (Liu et al., 2008)) and rain falls (2777 mm of water measured during the Morakot typhoon in 2009 (Ge et al., 2010)). Typhoons are mainly generated in either the South China Sea or the North-West Pacific Ocean East of the Philippine islands (Ou et al., 2002). From April to May, the spring winds correspond to a transitional period between monsoon and summer regimes.

Forced by the wind regimes described above, the hydrodynamics of the Taiwan strait (Jan et al., 2002) is strongly enhanced by the shallowness. The mean current is mainly southward during winter and northward during summer (reaching 1.7 m/s during the Haitan typhoon in 2005) (Chang et al., 2010, 2014). Typhoons have been observed to occasionally reverse the direction (Wu and Hsin, 2005; Zhang et al., 2013). Along the southwestern coast of Taiwan studied here, the tide is semi-diurnal and ranges about 2.68 m (Chen and Kuo, 2008).

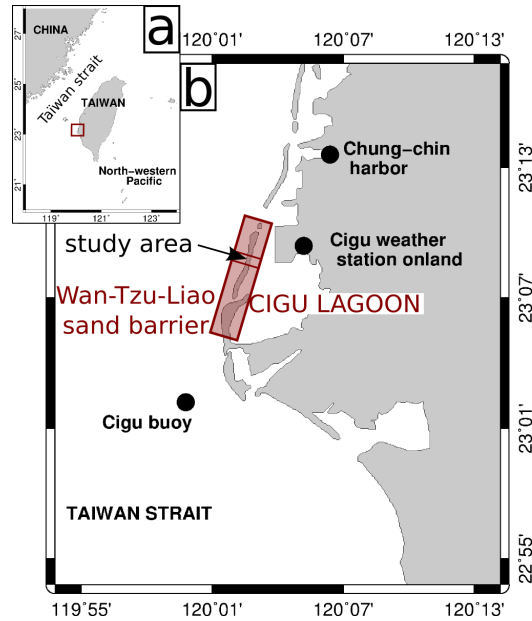
Chen and Kuo (2008) prove the eroding context of the region working on the Waisanding sand barrier 40 km northern the study area. Due to coastal structures interrupting sand supply from the major river Choushi (and others smaller rivers), the Waisanding sand barrier migrates south-westernward about 70-90 m/yr since 1986 coupled with erosion about 65 to 80 m/yr.

### The study area

The Cigu lagoon is one of the most attractive oyster-farming site along the south-western coasts of Taiwan. The lagoon is 3 km wide, 9 km long and 5 m of maximum water depth. The sediment is mainly sand transported within the lagoon by inlet currents and dune surge processes during storms. The lagoon shores are strongly anthropized. By contrast, the Wan-Tzu-Liao sand barrier which separates the lagoon from the open sea remains mainly a natural site (except extreme north and south parts where roads attach the barrier to the land (Campmas et al., 2014)).

The Wan-Tzu-Liao sand barrier is 9 km long. The width of the emerged barrier ranges between 100 m and 600 m with dune height between 1 and 10 m above the mean sea level (MSL). Occurrence of several wash-over fans and inlets in the historical archives gives evidence of constant dynamics of the barrier island through time. The long-term shoreline changes since 1993 suggest cross-shore sediments transfers in the north part. The elongated spit in the south area evidences a more important and regular

longshore transport of sediments driven by northern wind of monsoon reshaping the barrier southward. The randomized distribution of the inlets in the north suggests a weaker longshore drift in this zone with respect to the crossshore processes. The study area is defined in the north part of the barrier (middle island, in the vicinity of a tidal inlet formed in 2008) where cross-shore processes are dominant during typhoons performing a major control of the short and long-term shoreline changes (Campmas et al., 2014). Studied area is 200 m long  $\times$  300 m width with 2 to 3 m of dune height above MSL covered by vegetation. The grain size is almost homogeneous with  $d_{50}$  about 0.19 mm composed by 97.64 % of sand and 2.36 % of silt.



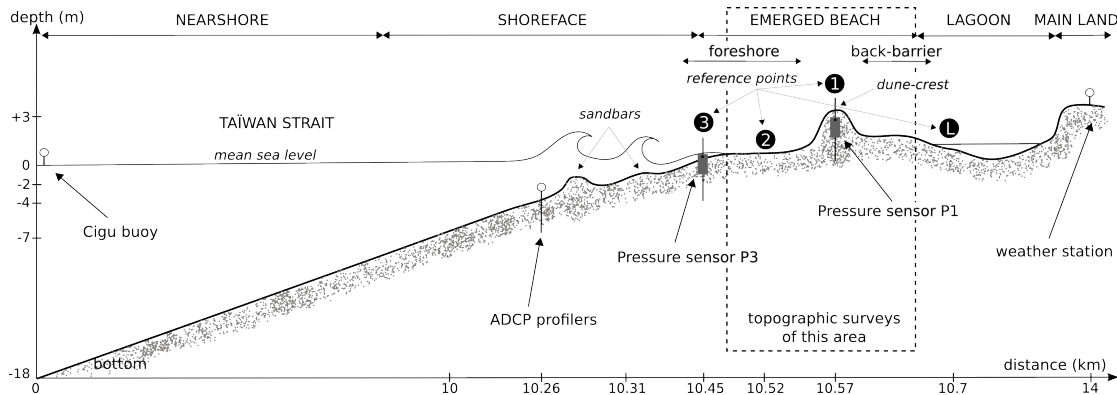
**Figure 1: Location of the study area. a) Taiwan in the North-western Pacific and b) the study area along the Wan-Tzu-Liao sand barrier system.**

## METHODOLOGY

In the framework of the KUN-SHEN project, 7 months of monitoring (from November 2011 to September 2012) provided offshore wind and wave, tide and surf to swash free surfaces measurements as well as topographic data of the emerged beach.

Offshore wind and wave conditions were obtained from the Cigu buoy (from *Taiwanese Central Weather Bureau*) located in 18 m of water depth. Wave conditions from surf zone to the swash zone were respectively acquired from one RDI current profiler (2Hz acquisition frequency) deployed in 4 m of water depth and pressure sensors (5Hz of measure frequency) deployed in the subtidal zone (P3) and buried in the dune crest (P1). Each instrument is fixed to a stainless steel frame (6 m long, 10 cm of diameter and 3 mm of thickness) deeply buried vertically in the sand (figure 2). In the lagoon, tide and inland wind conditions were respectively obtained from tidal gauge located northern in the Chung-Chin Harbor (figure 1b) and from inland weather station in the east coast (figure 2a).

Additionally, 20 topographic surveys of the emerged beach (i.e from the subtidal zone to the back-barrier) were conducted to provide Digital Elevation Models (DEM) of the beach, crossshore and longshore alti-bathymetric profiles and description of grain size distribution / composition. Surveys were performed each week in winter and just before and after each storm during summer. Using a kinematic RTK D-GPS (Trimble R8; +/- 5 cm in elevation and +/- 2 cm in X and Y), profiles are measured at the same cross-shore location for each survey while measurements of whole study area follow a grid at the accuracy of several points per  $m^2$ . Spatial resolution is maximal for areas of interest and breaking slopes (dune, cliff, sand bar, crescentic dune, washover fan).



**Figure 2: Schematic crossshore profile of the study area coupled with instruments location and typology used.**

All the DEMs (figure 3b or c), profiles (figure 3h) and sand volume changes (DEMs differential; figure 3d) maps were performed using Bash script and Generic Mapping Tools (GMT, (Wessel and Smith, 1998)). We use the Akima interpolation (based on spline functions providing smoother and much more natural curvatures, (Akima, 1970)) to build the DEMs. Isobaths and reference points (points number 1, 2, 3 and L) are plotted in order to improve the reading. Isobaths  $Z = 0$  m and 1 m above MSL are boundaries of subtidal (below 0 m of elevation), intertidal and supratidal (above 1 m of elevation) zones. Because of the lack of continuous morphological marker along the barrier (like dune-foot, dune-crest or berm), we defined a shoreline at the exact location of the isobath  $Z = 1$  m. On the sand volume changes maps, the solid and dashed lines display respectively isobaths at the first and last dates and the blank in the legend masks the morphological changes include into the margin of error.

Analysis of the morphodynamic changes driven by the Talim typhoon is based on two surveys : June 16<sup>th</sup> and June 22<sup>th</sup>; and seasonal morphodynamic changes are based on three surveys : October 21<sup>th</sup>, 2011, May 3<sup>th</sup>, 2012 and August 21<sup>th</sup>, 2012.

The empirical Bagnold equation (Bagnold, 1941) is used to estimate the potential aeolian sediment transport driven by strong wind of monsoon:

$$q = C(d/D)^{1/2} \cdot (\rho/g) \cdot U_*^3 \quad (1)$$

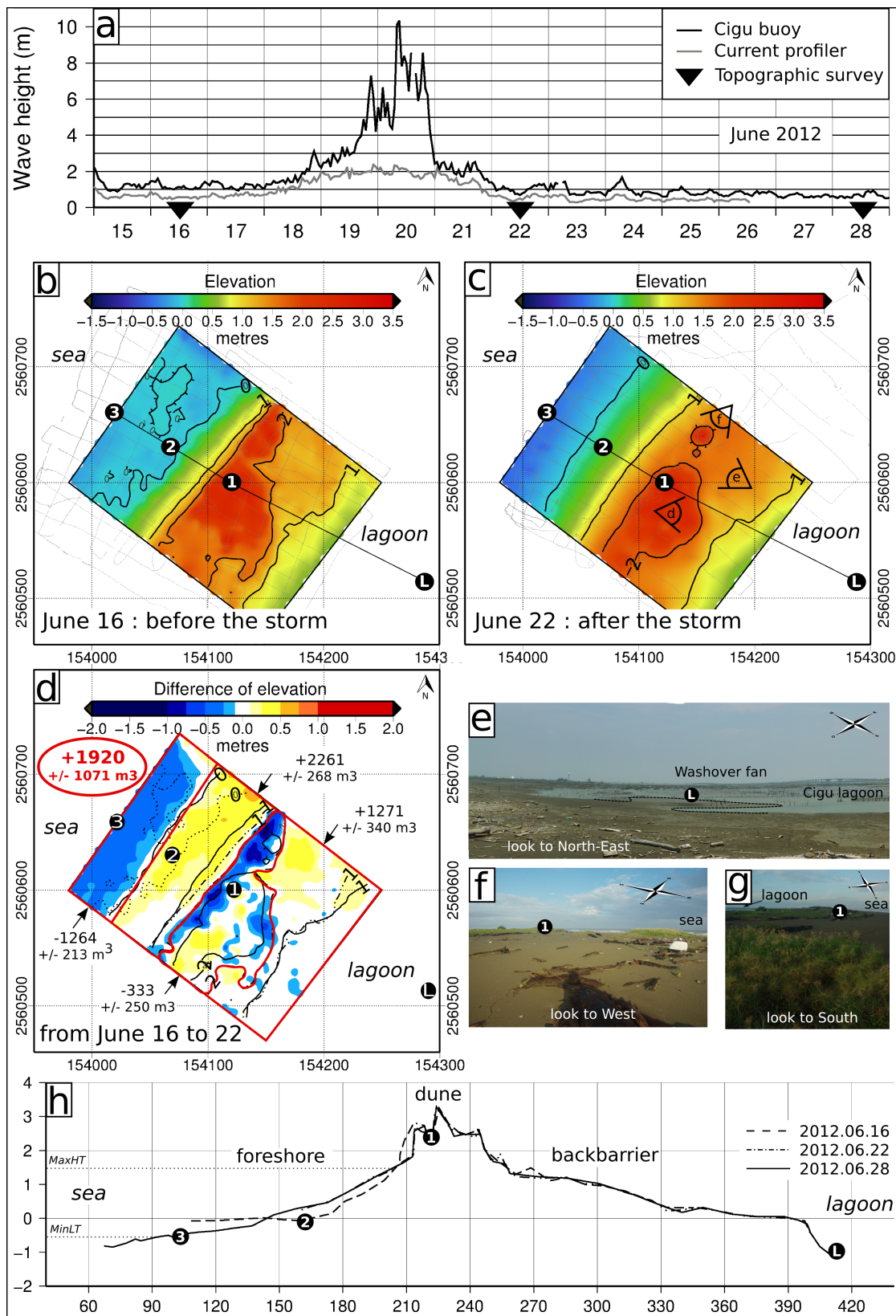
where  $q$  is the sand transport rate (the sand flux per unit time and per unit width, in kg/m/s), the coefficient  $C$  is 1.8,  $d$  is the grain diameter (m),  $D$  is a standard grain size diameter (0.00025 m),  $U_*$  is the shear velocity (m/s),  $\rho$  is the air density (1.22 kg/m<sup>3</sup>) and  $g$  is the gravitational force (9.81 m/s<sup>2</sup>).

## RESULTS

### Hydro-morphodynamics patterns driven by Talim tropical storm

Amongst the eight typhoons recorded, Talim is the most energetic observed. The storm, generated on June 18<sup>th</sup> in the east of China sea, travelled north-eastward to cross the study area on June 19-20<sup>th</sup>. The depression finally weakens on the Japan coasts on June 21<sup>th</sup>. We defined the storm apex at the maximum offshore wave height on June 20<sup>th</sup> between 7 and 9 am. At the storm apex, significant wave height reached 10.34 m offshore ( $T_p = 14.6$  s) and 2.3 m (period reaches 13.4 s) in the surf zone (figure 3 a). Wave periods measured in the surf zone are longer during the rising storm (16 s on June 19<sup>th</sup> 4 pm and 15.5 s on June 20<sup>th</sup> 2 am) than the apex or falling storm. At the shore, the wave set-up remains weak and most of the incoming energy is transferred into infragravity frequency band (Sous et al., 2013). The dune crest is not submerged on the studied transect even at the wave peak but morphological evidences of overtopping have been observed where the dune crest is lower (elevation about 2 m above MSL). At the storm apex, the wind is about 17.5 m.s<sup>-1</sup> offshore and 12 m.s<sup>-1</sup> inland (wind gusts about 20 m.s<sup>-1</sup>).

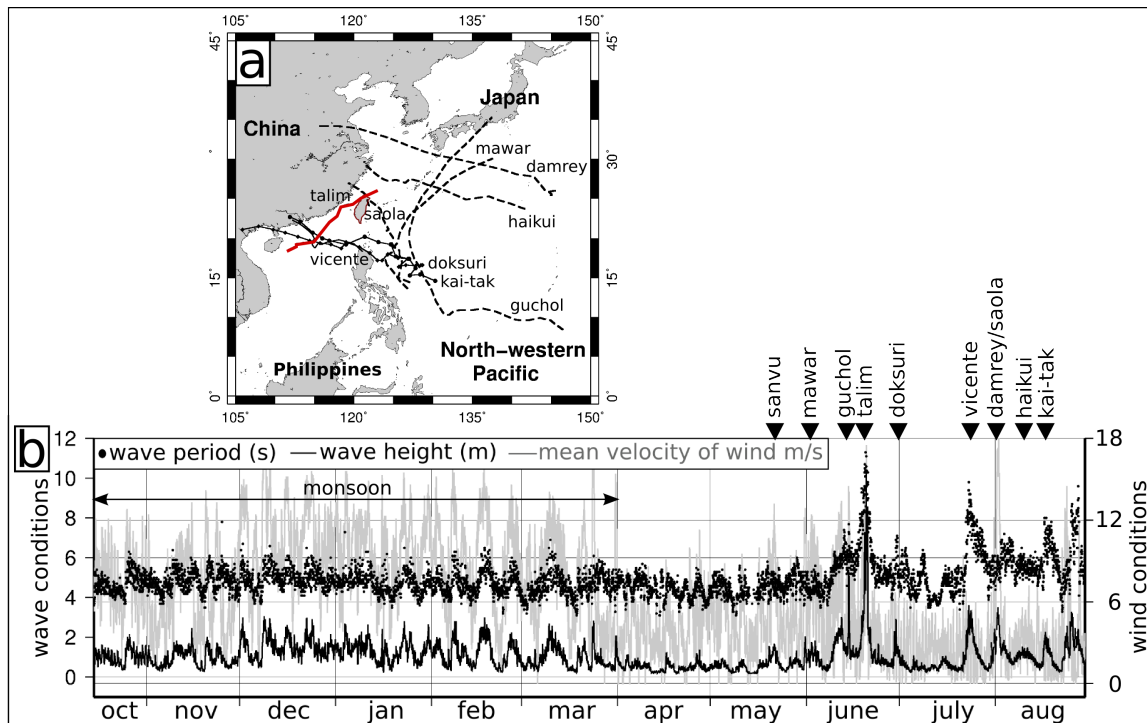
Regarding morphology, Talim tropical storm caused both expected and more unusual changes on the emerged beach. The expected changes include a 6.7 m dune-foot retreat (dune erosion about -333 m<sup>3</sup> +/- 250 m<sup>3</sup>) and 20 m long breach at the weakness point of the dune elevation (only 2.2 m above MSL). The sand transferred from the dune-breach to the back-barrier associated to wash-over deposit in the lagoon. These changes display a well-known morphodynamic responses to moderate storms but, more surprisingly, Talim drove also a significant nourishment of the upper-foreshore (+2261 m<sup>3</sup> +/- 268 m<sup>3</sup>) as well as the whole sand barrier (+1920 m<sup>3</sup> +/- 1071 m<sup>3</sup>) (figure 3 d). This sand input on the emerged beach results in 8 m of shoreline seaward shift and a change of the profile shape from concave to planar.



**Figure 3: Hydro-morphodynamics driven by Talim tropical storm. a) Offshore and surf zone waves conditions during the storm. DEMs on b) June 16<sup>th</sup> and c) June 22<sup>th</sup>, used to characterize the d) morphologic changes of the emerged beach driven by the typhoon. The black line displays the crossshore profile plotted on h). Orientation of camera pictures e, f and g) are display on map b).**

### Hydro-morphodynamics patterns driven the typhoons and the monsoon seasons

During fair-weather periods separating each typhoon of the summer season, offshore wave conditions remain calm, with  $H_s < 1$  m and  $T_p < 5$  s. During typhoons, waves ranges from  $1 < H_s \leq 10$  m with  $5 < T_p \leq 14$  s offshore. Strong wave dissipation and spectral energy transfers occurs over an expectedly very wide surf zone, as indicated by the smaller wave height range  $1 < H_s \leq 2.3$  m measured by the profiler. During winter, regular moderate storms induced by Asia wind of monsoon reaches occasionally 3 m of significant offshore waves. The period remains below 7 s (6.4 s in average). The constant North wind velocity is about 7.6 m/s in average resulting in a significant aeolian sand transport on the emerged beach. Order of magnitude of the potential sand transport computed with classic Bagnold equation is about 0.23  $\text{m}^3/\text{h}$  along the season.



**Figure 4: Wind and wave conditions along the year 2011/2012. a) Tracks of typhoons monitored. b) Offshore wind and waves along the monsoon and typhoons seasons. Wind from June 21<sup>th</sup> to August were extracted from the inland weather station (figure 1 c and 2).**

Surprisingly, the summer results in an accretion period with a sand input about  $+3556 \text{ m}^3 \pm 912 \text{ m}^3$  mainly transferred to the foreshore inducing a 10.44 m shoreline seaward shift (figure 5 b,f). The sand balance evidences a sand transport probably from the surf zone and/or the shoreface. Along the season, energetic conditions during moderate typhoons (offshore significant wave below 3 m) drove the sandbars to aggrade upward from the subtidal to the supratidal zone (figure 5g). Extreme energetic event (offshore significant wave above 5 m) like Guchol (from 5<sup>th</sup> to 16<sup>th</sup> June) or Talim induces the sandbars destruction and the sand-stock redistribution over the foreshore (see the landward migration of sandbars from May 3<sup>th</sup> to the Talim event on the figure 5).

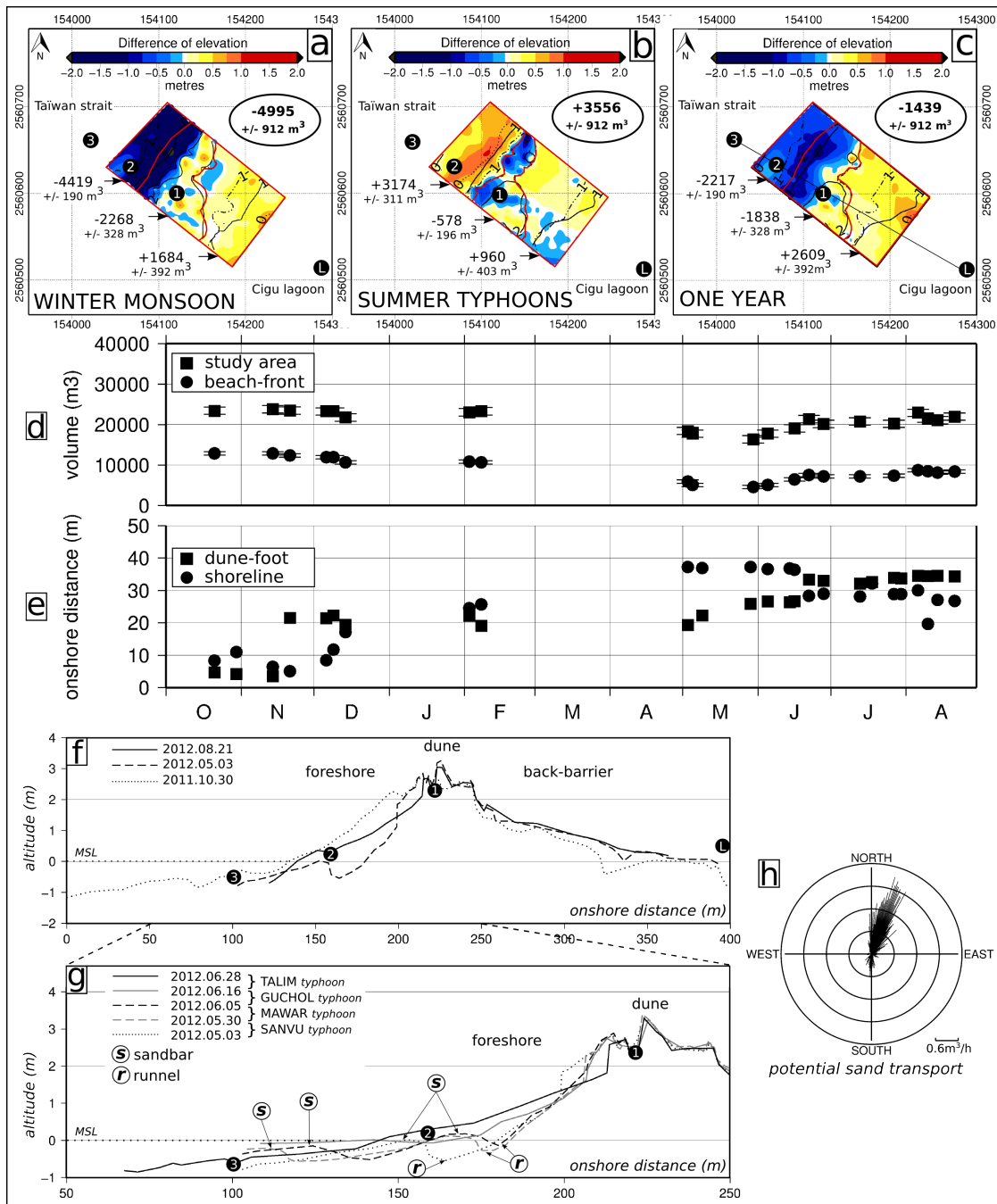
The dune is eroded ( $-578 \text{ m}^3 \pm 196 \text{ m}^3$ ) and retreats landward (15.06 m). The back-barrier nourishment ( $+960 \text{ m}^3 \pm 403 \text{ m}^3$ ) is allocated to the cross-shore processes because heavy rain prevents aeolian sand transport during typhoon events.

Conversely, winter are erosive season with negative sand balance about  $-4995 \text{ m}^3 \pm 912 \text{ m}^3$  (figure 5 a,f). The regular collision regime (Sallenger, 2000) driven by moderate storms induces foreshore erosion about  $-4419 \text{ m}^3 \pm 912 \text{ m}^3$  resulting in 28.9 m of shoreline retreat, dune scarp (erosion about  $-2268 \text{ m}^3 \pm 328 \text{ m}^3$ ) and 28 m of dune-foot retreat. Aeolian sand transport occurs on the dune and the back-barrier

(figure 5 h) but no significant surge process of the dune occurs during this season.

Over the year, the sand barrier recorded 18.5 m of landward shift coupled with dune erosion ( $-1838 \text{ m}^3 \pm 328 \text{ m}^3$ ) and dune-foot landward retreat (29.6 m) without crest abrasion (figure 5 c,f). The back-barrier nourishment ( $+2609 \text{ m}^3 \pm 392 \text{ m}^3$ ) is a mix between aeolian sand input during North wind of winter and sand transport from beach-front during typhoons. The negative sand balance about  $-1439 \text{ m}^3 \pm 912 \text{ m}^3$  suggests a sand transfer outside the study area.





**Figure 5: Morphodynamic patterns driven by the monsoon and typhoons seasons. DEMs on October 30<sup>th</sup> 2011, May 3<sup>th</sup> and August 21<sup>th</sup> 2012 are used to build sand volume changes maps of the a) monsoon and b) typhoons seasons as well as c) the whole year. Solid and dashed lines display respectively isobaths at the youngest and oldest dates and the sand balances correspond to the respective red frames. The straight black line displays the crossshore profile plotted on f and g). The sand volume changes of the study area/beach-front and location changes of the shoreline/dune-foot along the year are plotted respectively on graphs d) and e). h) Rose of the potential aeolian sand transport during the monsoon season**

**DISCUSSION**

**TALIM tropical storm drives a foreshore nourishment**

Only few researches have demonstrated a subaerial beach accretion after an extreme storm passage. Tapia et al. (2014) present in the ICCE conference 2014 (The 34th International Conference on Coastal

Engineering) a subaerial beach accretion in Puerto Morelos (Mexico) after the Wilma hurricane in October 2005. Roberts et al. (2013) also measured 0.7 m of thick layer of sand on the berm crest prograding the beach seaward driven by the hurricane Bill in 2009, Florida. Stone et al. (2004) also showed that barrier islands can conserve mass during catastrophic cyclones, and that less severe cyclones and tropical storms can promote rapid dune aggradation and can distribute sediment to the entire barrier system.

In our case, we measure an unexpected nourishment of the whole emerged beach during Talim tropical storm which is more classically observed in the literature under beach recovery process conducted by fairweather conditions (Komar, 1998; Larson and Kraus, 1994; Ruggiero et al., 2005) or inter-hurricanes periods (Roberts et al., 2013). Because the beach-front appears mainly controlled by cross-shore processes during summer, the emerged beach nourishment probably results from a landward shift of sand coming from the shoreface.

#### **The beach recovery during typhoon season is related to the landward migration of the intertidal sandbars.**

Quite surprisingly, the sand balance displays an accretive period of beach-front recovery and shoreline seaward shift during summer. The sand input appears occurring through a landward migration and upward aggrading sandbars in the intertidal zone that served as the sediment path-way for the emerged beach reconstruction driven by energetic conditions during typhoons.

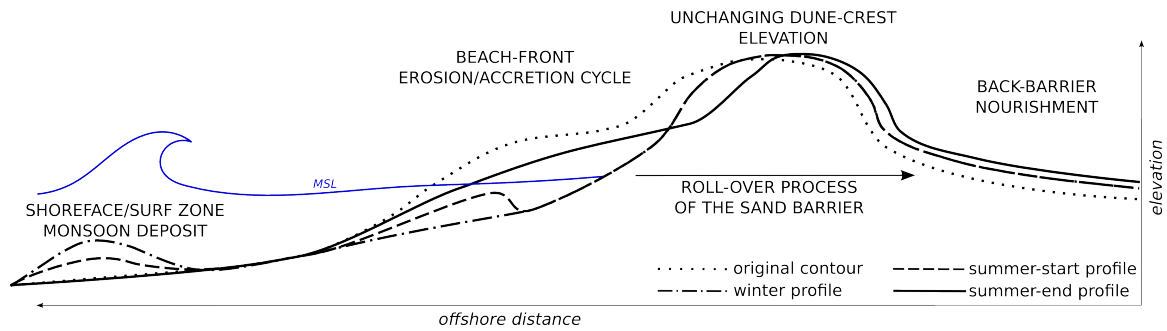
Similar observations of intertidal sandbar destruction during storm were reported by Aagaard et al. (1998). Masselink et al. (2011) suggest that the intertidal sandbar represents a small-scale version of a transgressive barrier system occurring through the four regimes described by Sallenger (2000) for the behaviour of barrier islands during hurricanes (swash, collision, overwash and inundation). In our case, we observed a landward migration of the intertidal sandbar driven by energetic, but non-destructive typhoons (offshore significant wave below 3 m), until destruction and sand redistribution over the foreshore driven by paroxysmal events (offshore significant wave above 5 m) like Guchol or Talim. In addition to the hydrodynamic conditions, the feedback with morphology can play a significant role in the sandbar dynamics. Current meter records obtained by Masselink et al. (2011) suggested that the onshore migration of an intertidal bar was driven by onshore-directed near-bed mean currents which persisted for as long as a runnel was present landward of the bar. When this runnel infilled, the mean current across the bar changed direction, becoming an offshore-directed bed return flow that destroyed the bar. This behaviour corresponds to our measures. The sandbars migrate landward until the runnel filling during Guchol event and the complete destruction during Talim event flattening the beach profile. But, the sandbar volume cannot explain the entire positive sand balance, that suggests a concomitant sand supply from the shoreface. Indeed, during paroxysmal events, the positive sand balance results from a mix between the sand redistribution from the intertidal sandbars and a direct shift of sand from the shoreface.

#### **The major role of the monsoon season in the beach-front erosion process**

The major agent of beach-front erosion are the moderate storms series during the monsoon season. Roberts et al. (2010) under hurricanes climate, and many others under moderate storms (Short, 1999), assigned a major role to the inter-storm periods in the reconstruction process of the emerged beach. Working on model along the Portuguese west coasts, Ferreira (2005) proved that storm groups with relatively small return periods induce average erosion volumes as significant as a single storm with a much longer return period. In this study, 20 moderate storms above 2 m of offshore significant wave height (i.e. an average of 3.3 storms per month) completely inhibit the inter-storms reconstruction process which induces an overall beach-front erosion. Because the alongshore drift is present during the monsoon, a part of the sand taken from the beach-front is probably lost in the drift when the major part should be deposited seaward across the shoreface.

#### **One conceptual model to display the yearly beach behaviour**

Yearly, the volume of sand transported in the shoreface during monsoon will be in average returned to the emerged beach driven by energetic conditions of typhoons. Consequently, the major importance of typhoons on the roll-over processes is to transport the sand landward to the foreshore and the back-barrier through dune surge processes (overwash and inundation regimes). Indeed, we propose a conceptual model to display the sand barrier morphodynamics driven by monsoon and typhoons conditions along one year (figure 6).



**Figure 6: Conceptual model of the sand patterns along one year and successive monsoon and typhoon seasons.**

### Decadal tendency

However, the 18 m of sand barrier landward shift measured in 2011/2012 is lesser than expected when studying the decadal shoreline changes along the Wan-Tzu-Liao sand barrier (Campmas et al., 2014). The shoreline changes measured in this study are about 25 m/yr from 1999 to 2004 and 48 m/yr from 2004 to 2009. Likewise, the small loss of sand in 2011/12 ( $-1439 \text{ m}^3 \pm 912 \text{ m}^3$ ) compared to the rapid decrease of the sand barrier area observed since 2004 (Campmas et al., 2014) tends to categorize the 2011/2012 studied period as a rather rare year with an unusual frequency of events like Talim driven accretive morphologic patterns.

If typhoons with accretive morphologic patterns are the major agent of reducing shoreline retreat, the occurrence of events like Talim can play a major role in the long-term barrier dynamics and seems to be variable from one year to another resulting an irregular historic shoreline changes.

### CONCLUSION

This study focuses on the Wan-tzu-liao sand barrier morphodynamics at different time scales ranging from the typhoon event to entire typhoons and monsoon seasons. The three main objectives were, (i), to provide detailed description and quantification of hydro-morphodynamic changes under a single typhoon and a typhoons/monsoon seasons, (ii), to better understand relationships between extreme forcings and sand barrier dynamics and, (iii), to discuss these in light of existing conceptual models mainly developed from moderate storms data.

Seven months of field work provided a robust dataset of hydrodynamics conditions from offshore to the swash zone coupled with a high frequency surveys of the emerged beach. Based on this dataset, we observed an accretive morphological pattern of typhoon (Talim tropical storm) and typhoons season (2012). Inversely during monsoon, the increased storm occurrence frequency inhibits the reconstruction beach process resulting in beach-front erosion. A new conceptual model of sand barrier yearly morphodynamics is proposed. It combines a seaward transfer of sand taken from the beach-front during monsoon and a landward return of sand in the foreshore and the back-barrier during summer. However, compared to the decadal analysis of the sand barrier landward migration showed by Campmas et al. (2014) study, the shoreline changes measured in 2011/12 should be considered as a rather rare year with an unusual occurrence of Talim storm-like events with beach-front recovery pattern.

### ACKNOWLEDGEMENTS

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