



36TH INTERNATIONAL CONFERENCE ON COASTAL ENGINEERING 2018

Baltimore, Maryland | July 30 – August 3, 2018

The State of the Art and Science of Coastal Engineering

Hindcasting of surge and wave on Hokkaido coasts by a winter low pressure system using surge-wave coupled sea bottom and surface stresses in SuWAT

Sooyoul Kim; Tottori University

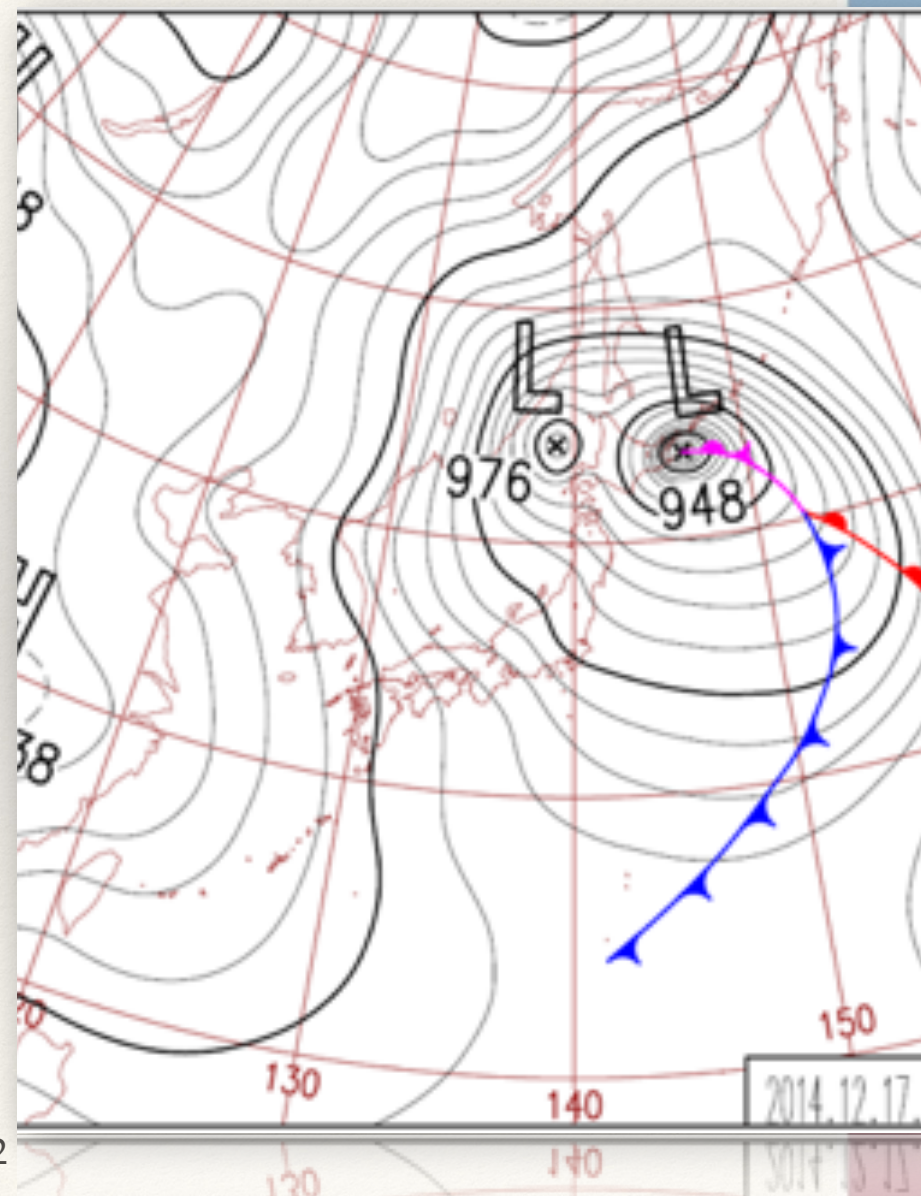
K., Kumagai, D., Tsujio and T., Tsuji; Pacific Consultants Co. Ltd.

Hajime Mase; Kyoto University



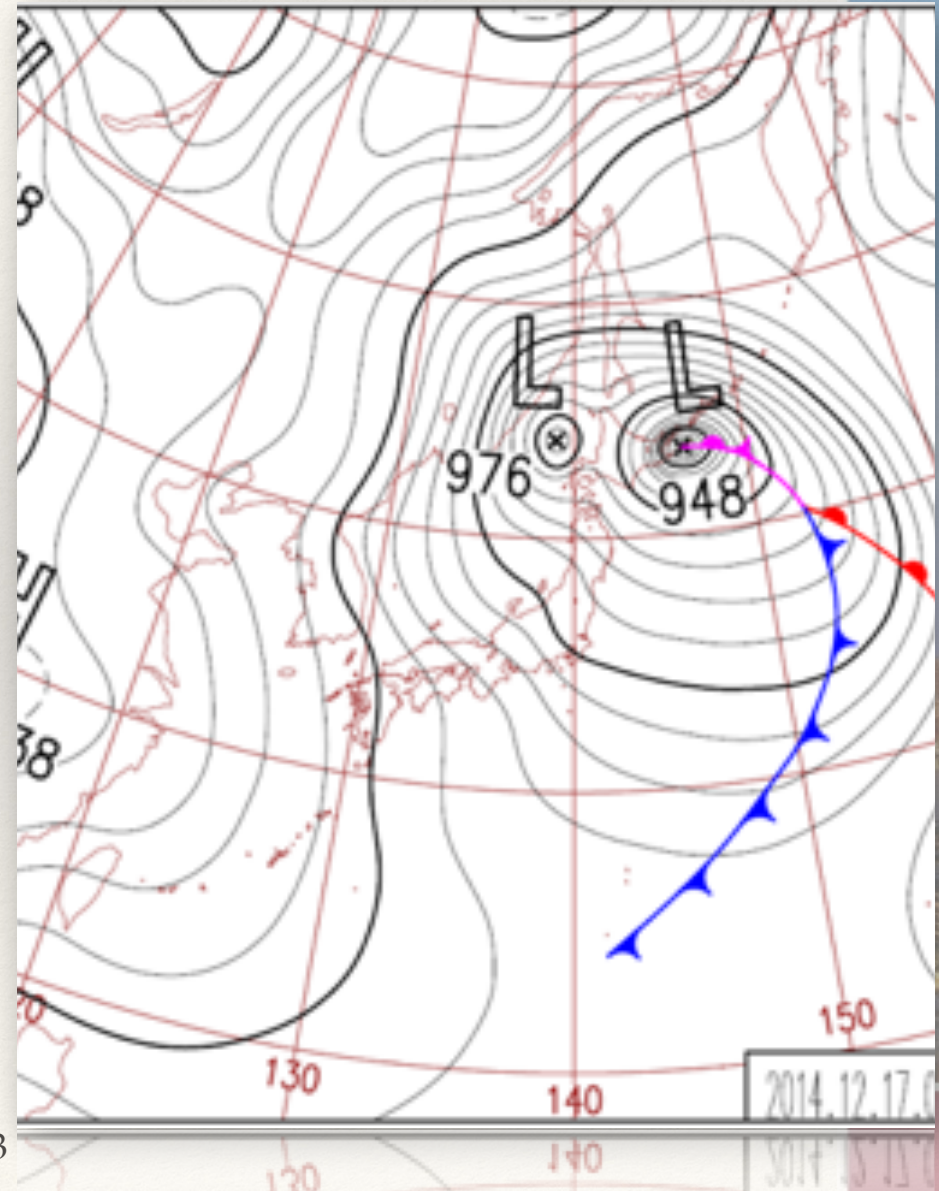
Background

- ❖ A Winter Low Pressure System (or Extratropical Cyclones) (WLPS): similar impacts to a typhoon.
- ❖ A WLPS In Hokkaido, Japan
 - ❖ 16-17 December 2014
 - ❖ 1.75 m high surge level
 - ❖ vast areas of flooding on the Kushiro coast
 - ❖ coastal facilities broken
 - ❖ New record-break surge level due to WLPS since record-break level of 0.9 m in 1994



Purpose

- ❖ Understanding the WLPS event in 2014
- ❖ Understanding surges and waves due to WLPS 2014
 - ❖ Hindcast: surges and waves
 - ❖ using a coupled model of surge, wave and tide
 - ❖ no consideration: tide

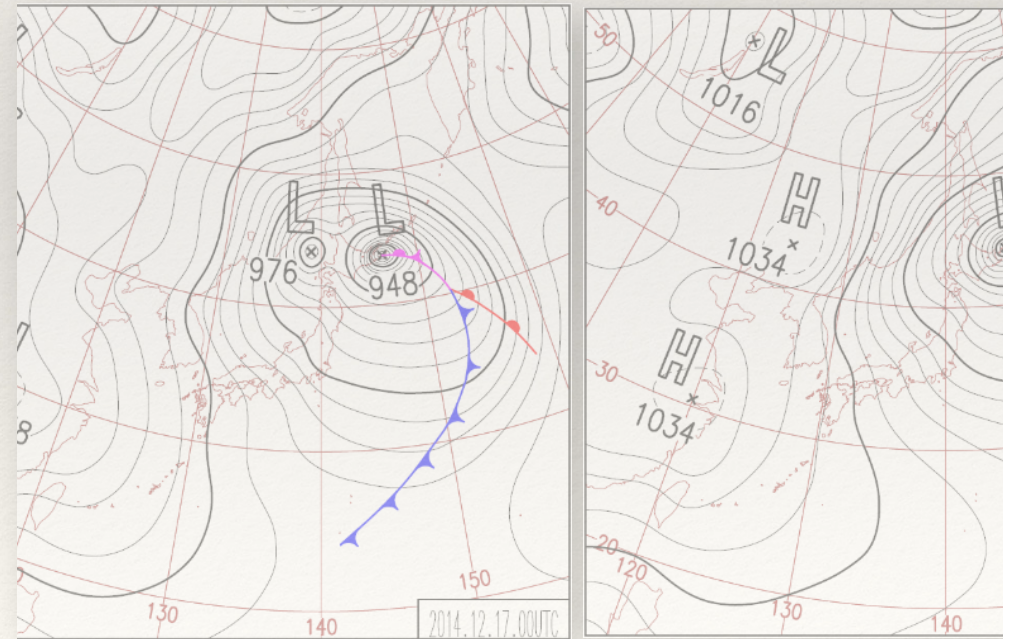
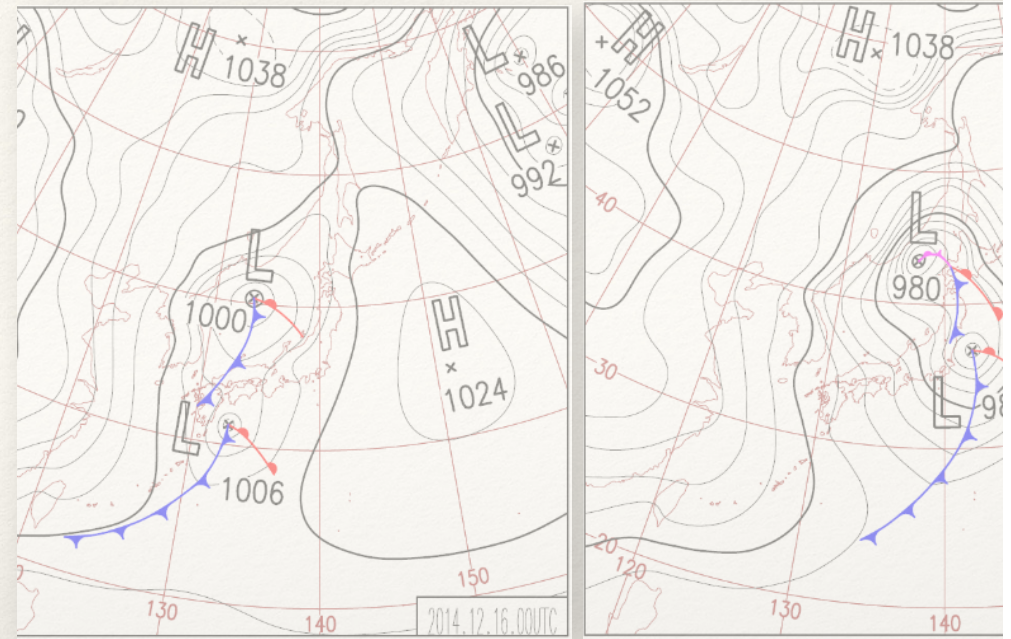




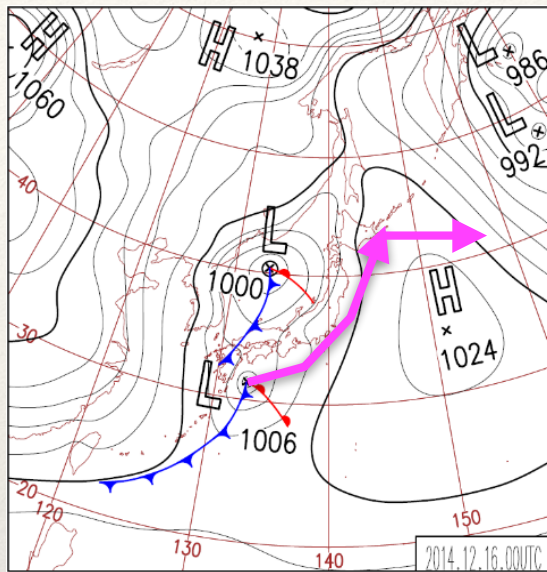
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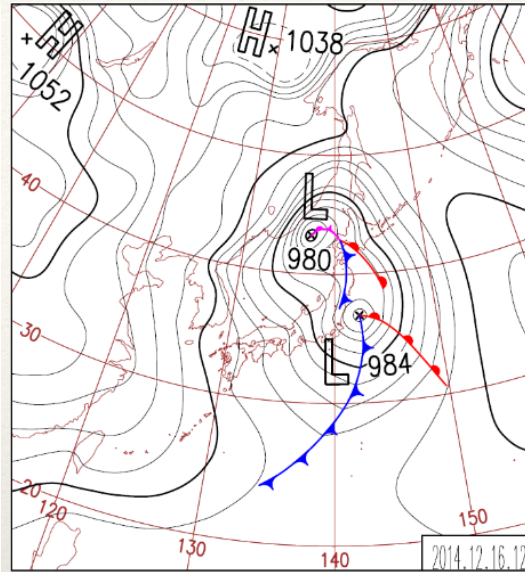
Winter Low Pressure System 2014



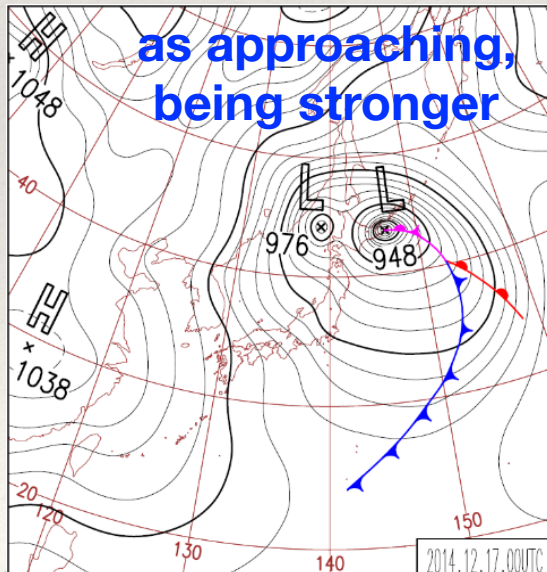
WLPS in Dec. 2014 vs Feb. 1994



9:00, 16th Dec. in JST

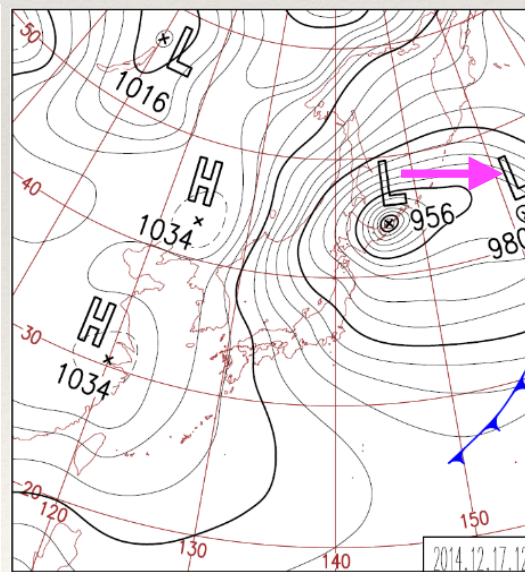


21:00, 16th Dec.

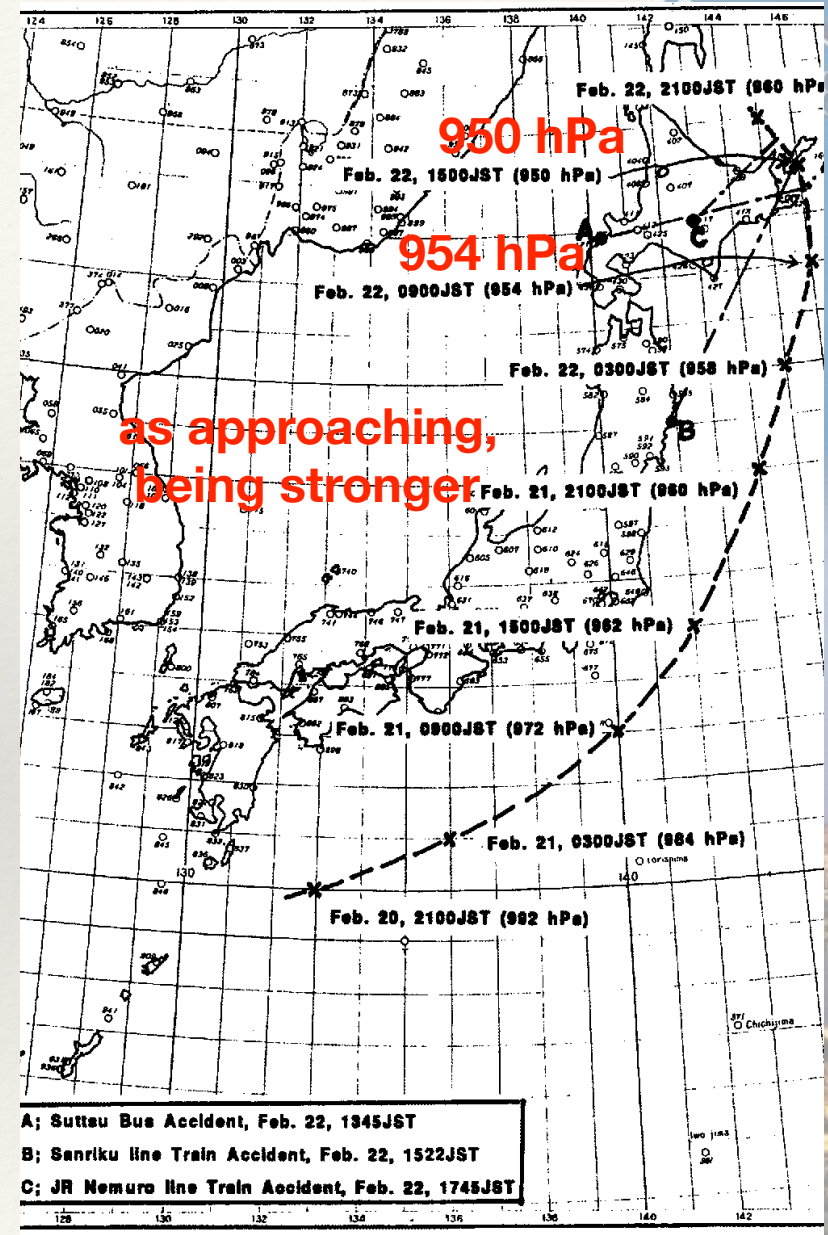


as approaching,
being stronger

9:00, 17th Dec.



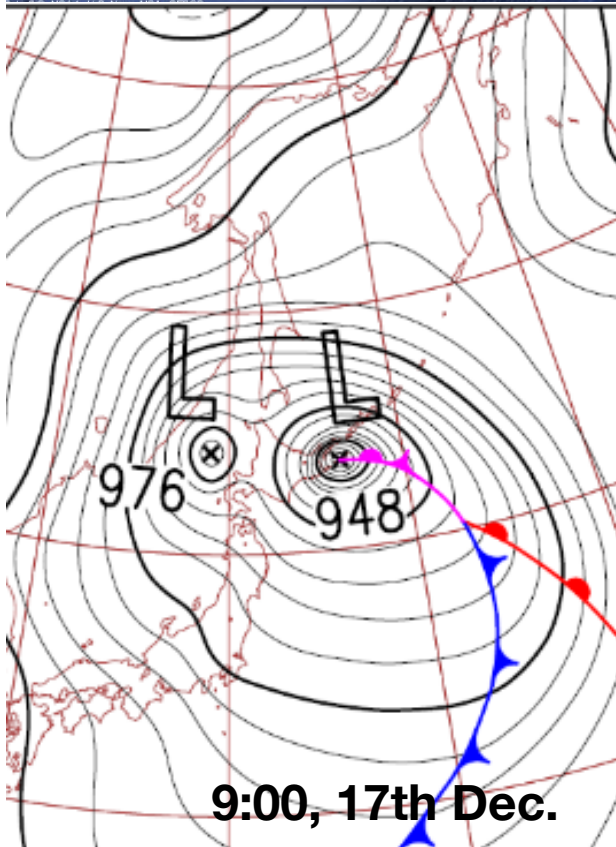
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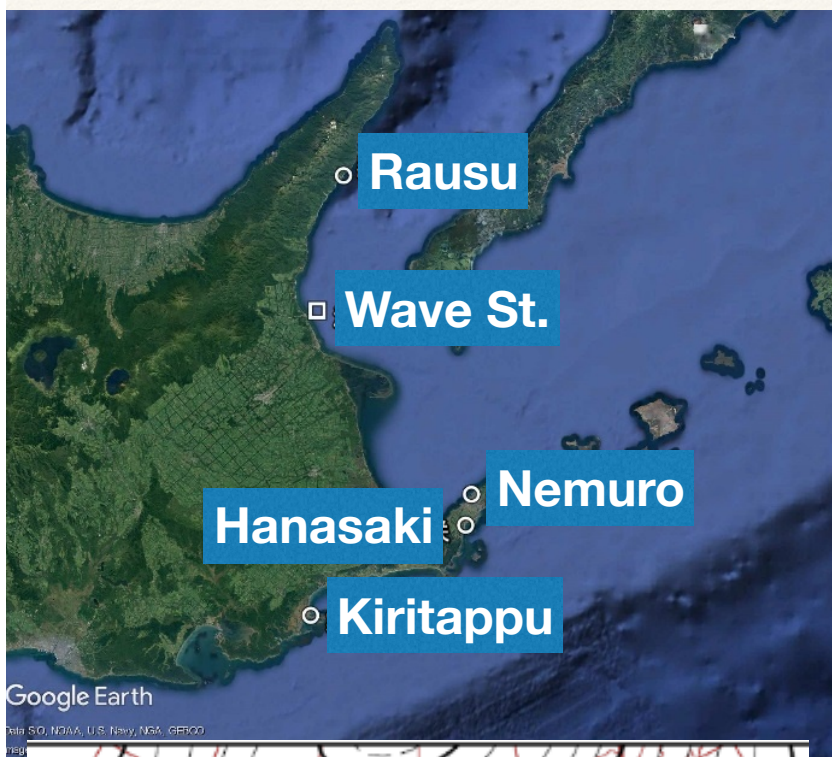


as approaching,
being stronger

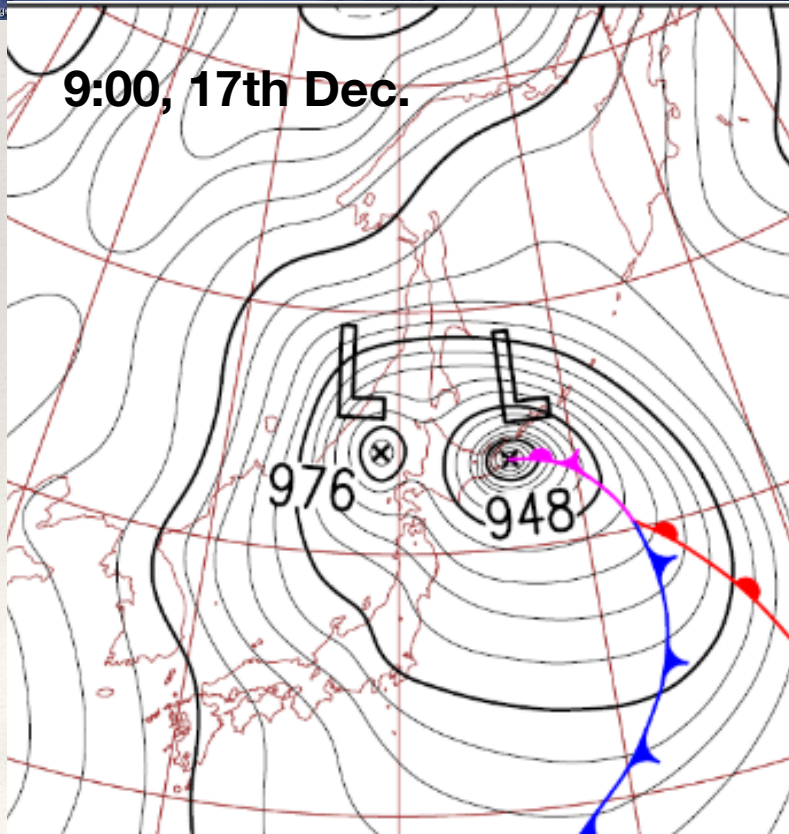
- A; Suttsu Bus Accident, Feb. 22, 1945JST
- B; Sanriku line Train Accident, Feb. 22, 1522JST
- C; JR Nemuro line Train Accident, Feb. 22, 1745JST

Observed wind and sea level pressure in Nemuro

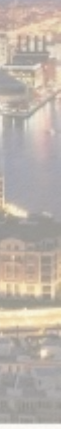
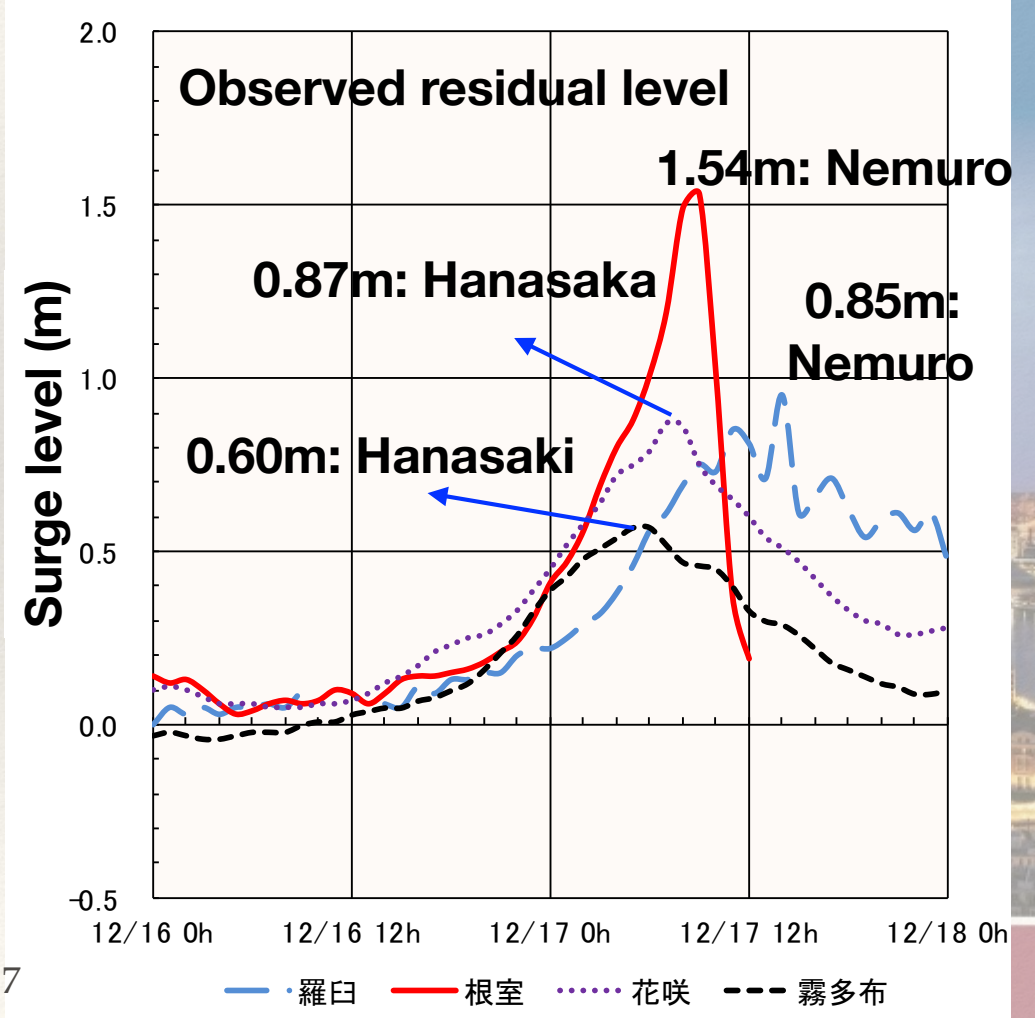
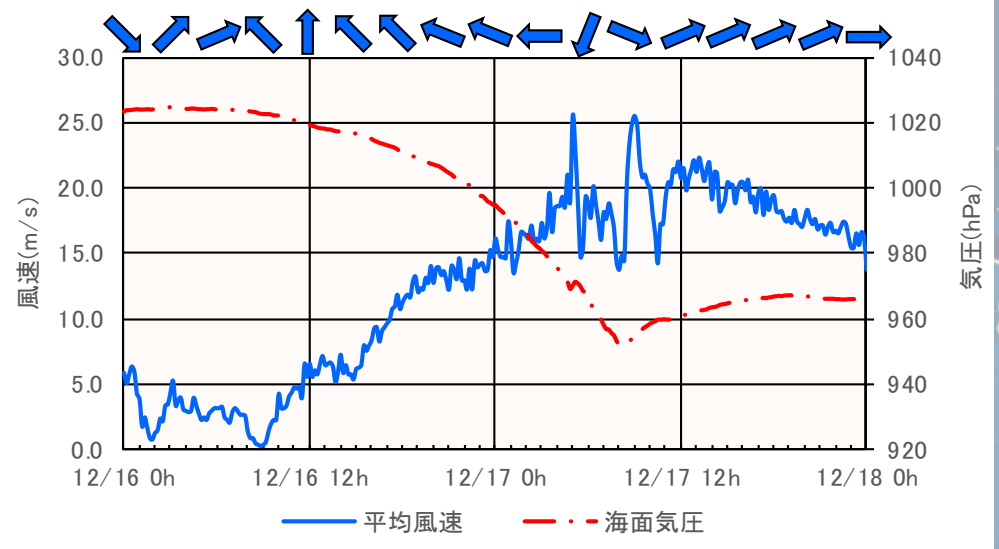


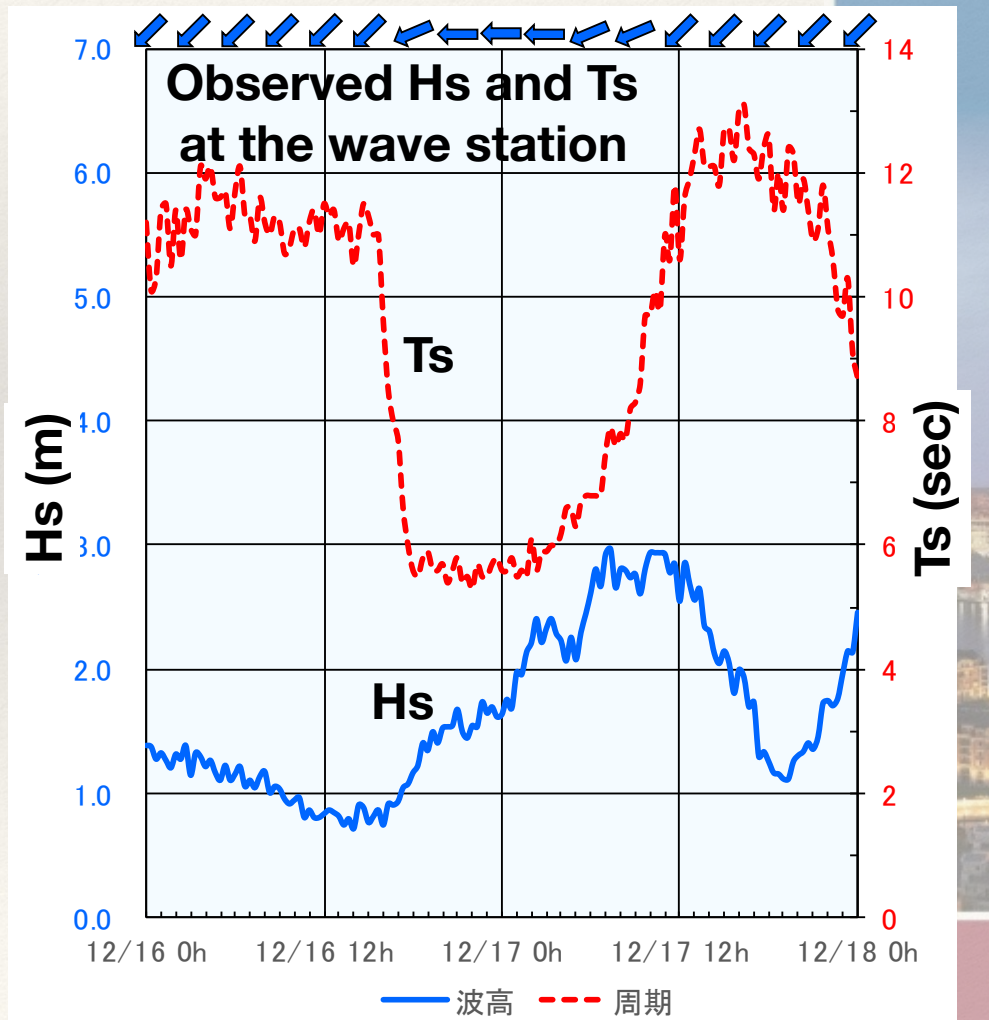
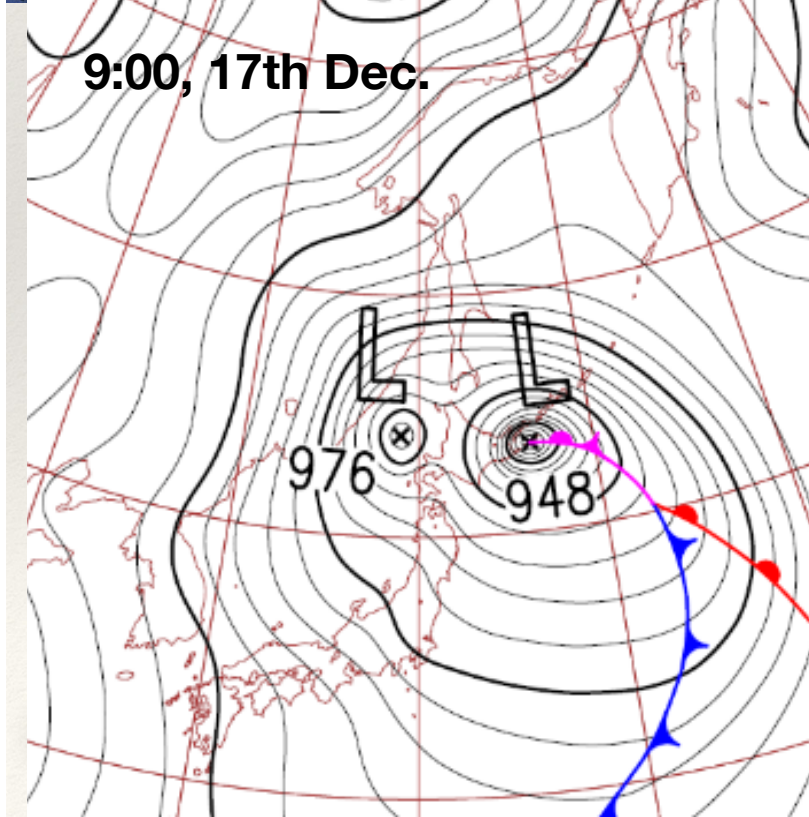
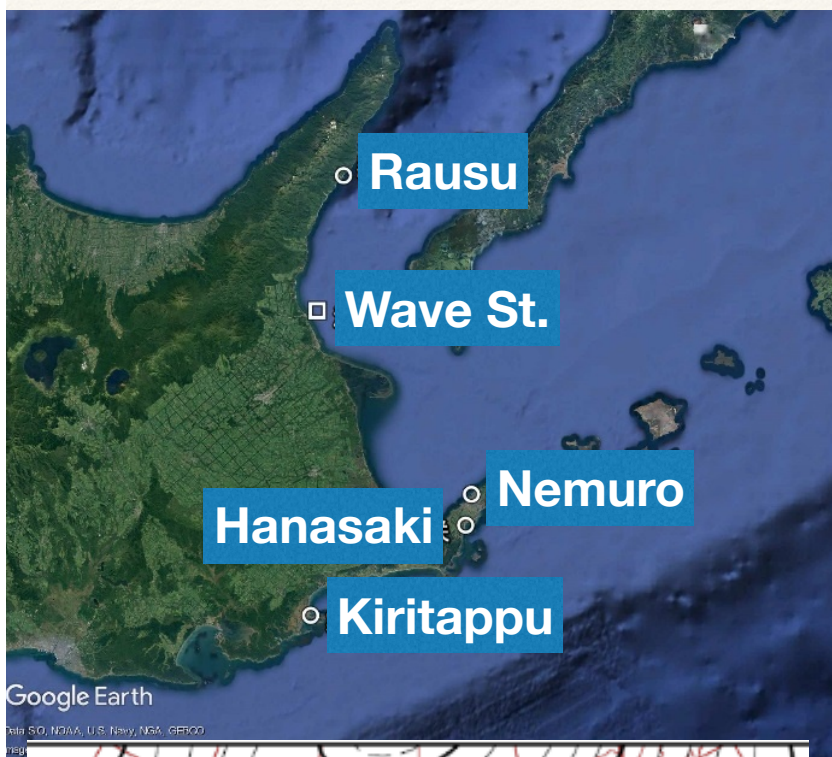


Google Earth
Data S.O, NOAA, U.S. Navy, NGA, GEBCO



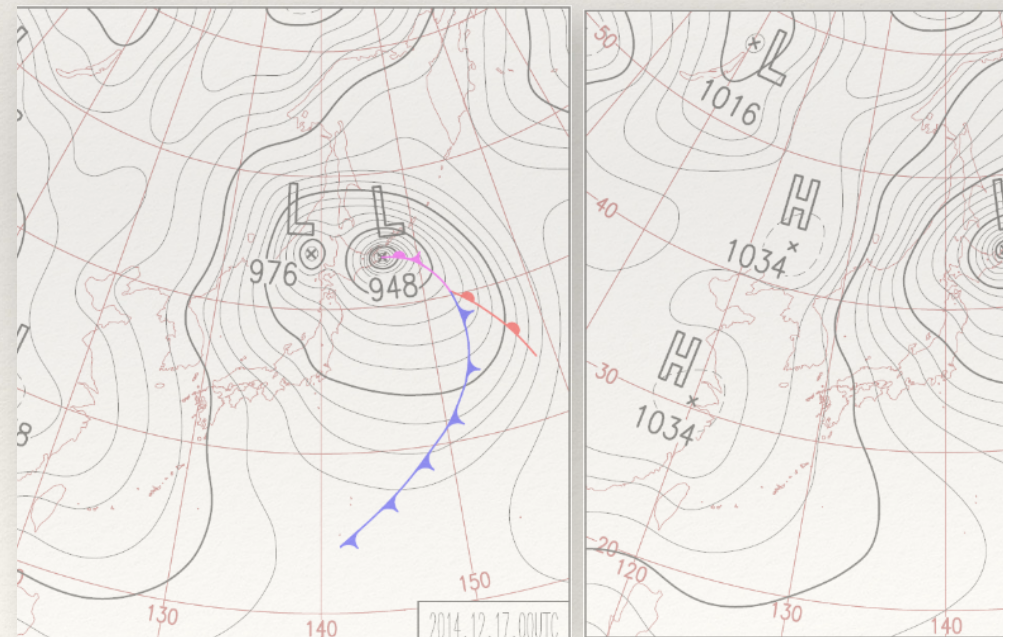
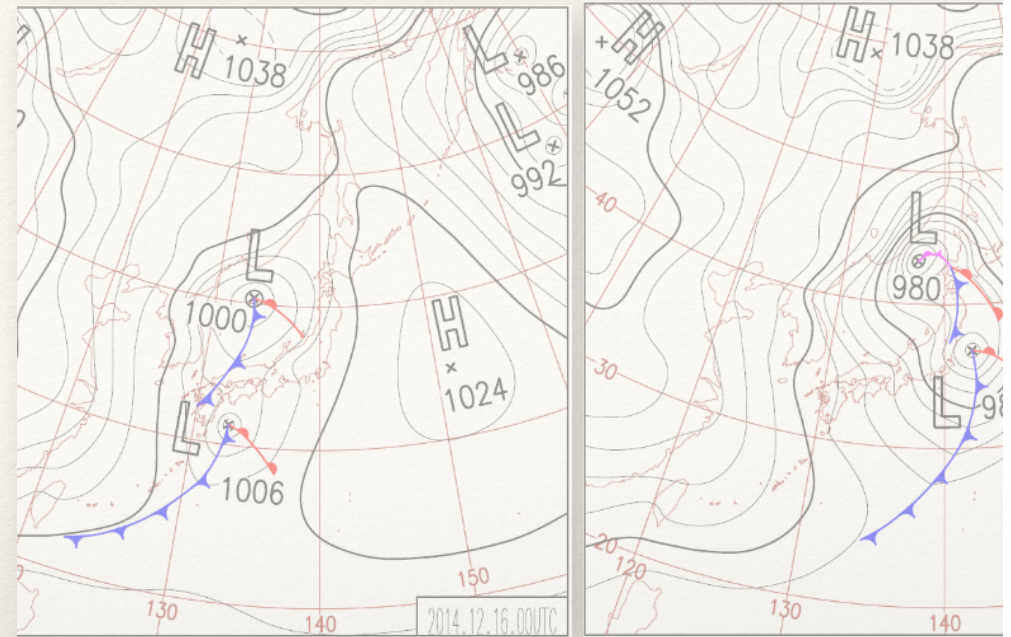
Consultants





Hindcasting surges and waves

- The coupled model of surge, wave and tide; SuWAT
- Wave dependent drag coefficient capped at specific wind speeds
- wave-current interaction-induced bottom stress

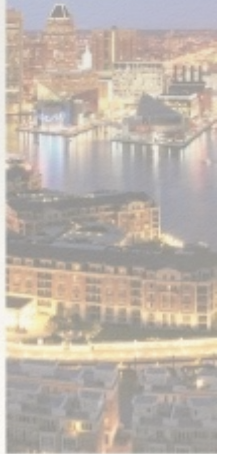


The drag coefficient, C_d , in sea surface layer

- ❖ Wave growth term in SWAN

$$S_{\text{in}}(\sigma, \theta) = A + BE(\sigma, \theta)$$

- ❖ A: the linear wave growth term
- ❖ BE: the exponential wave growth term



The drag coefficient, C_d , in sea surface layer

- ❖ Wave growth term in SWAN

$$S_{\text{in}}(\sigma, \theta) = A + BE(\sigma, \theta)$$

- ❖ A: the linear wave growth term
- ❖ BE: the exponential wave growth term
- ❖ C_d in the linear wave growth term
 - ❖ Transfer U_{10} to U^* the friction velocity ($u_*^2 = C_D U_{10}^2$)

Wu (1982):

$$C_D = \begin{cases} 1.2875 \times 10^{-3} & \text{for } U_{10} < 7.5 \text{ m/s} \\ (0.8 + 0.065 U_{10}) \times 10^{-3} & \text{for } U_{10} > 7.5 \text{ m/s} \end{cases}$$

Zijlema et al (2012):

$$C_D = (0.55 + 2.97\tilde{U} - 1.49\tilde{U}^2) \times 10^{-3}$$

The drag coefficient, C_d , in sea surface layer

- ❖ Wave growth term in SWAN

$$S_{in}(\sigma, \theta) = A + BE(\sigma, \theta)$$

- ❖ BE: the exponential wave growth term
- ❖ Janssen's wave dependent C_d in the exponential wave growth term (1991) and following Mastenbroek et al.(1993) accounting for sea state

- ❖ Wind profile: $U(z) = \frac{u_*}{\kappa} \ln \left(\frac{z + z_e + z_0}{z_e} \right)$ Turbulent stress $\tau_t = \rho_a (\kappa z)^2 \left(\frac{\partial U}{\partial z} \right)^2$

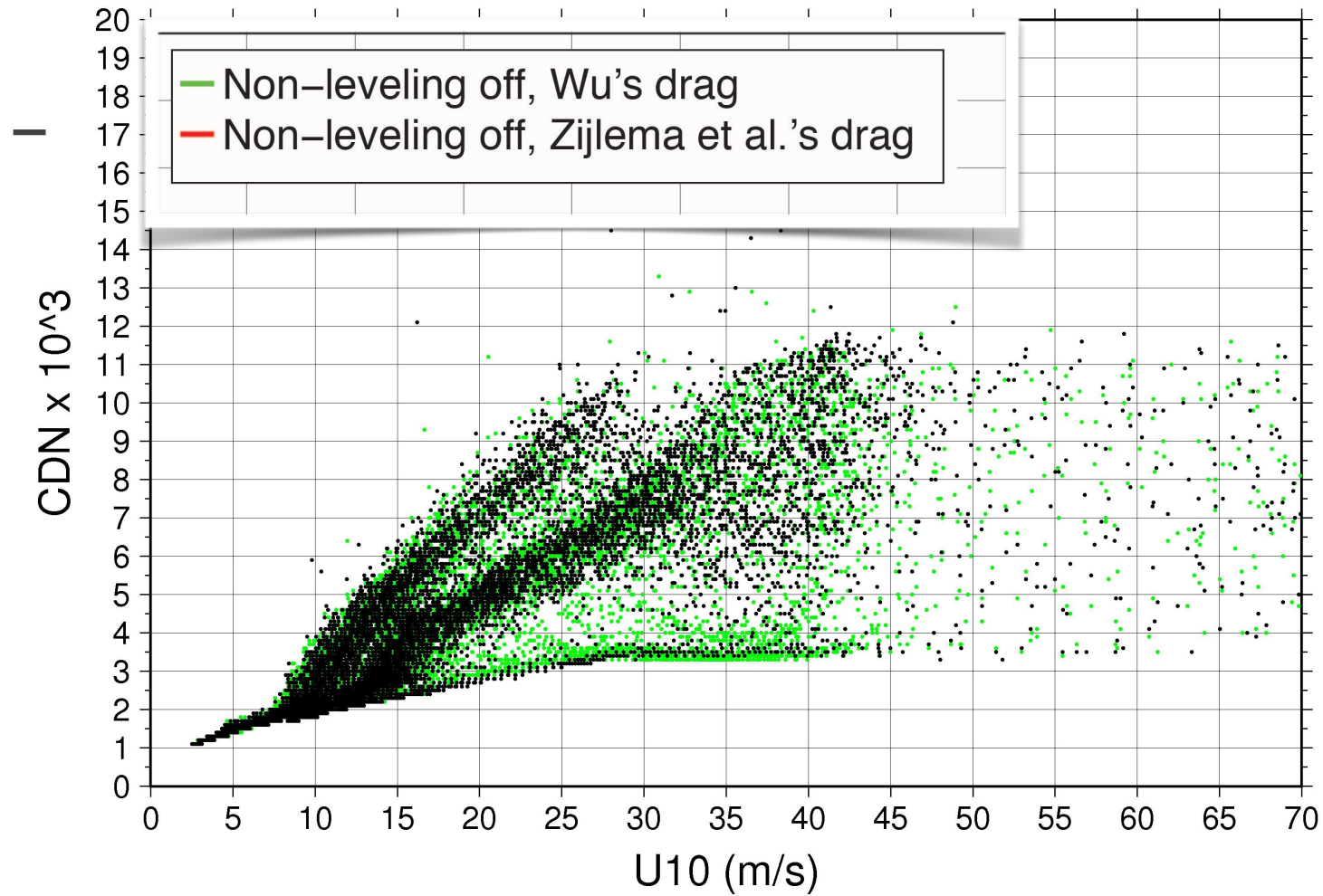
- ❖ Effective roughness: $z_e = \frac{z_0}{\sqrt{1 - \tau_w/\tau}}$

- ❖ Wind speed-capped Wave dependent C_d

$$C_D = u_*^2 / U(z)^2 = \left[\kappa / \ln \left(\frac{z + z_e - z_0}{z_e} \right) \right]^2$$

$$\tau_s = \rho_a C_D \vec{U}_{10} |\vec{U}_{10}|$$

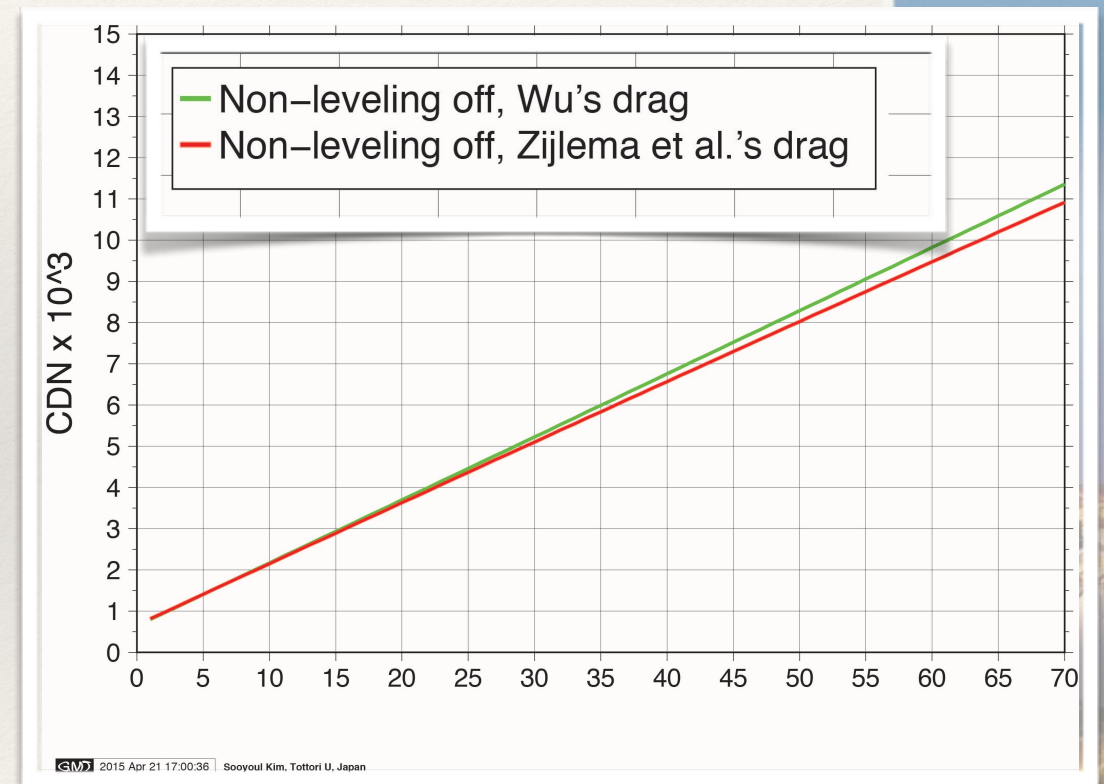
Estimated wave dependent C_d without levelling off due to Typhoon Hainan 2013



The best-fitted wave dependent C_d to the 2nd-order polynomial due to Typhoon Hainan 2013

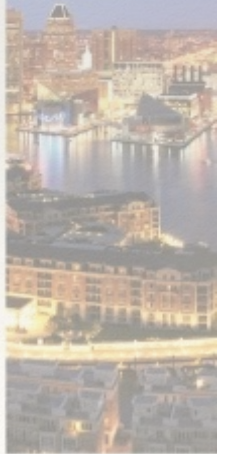
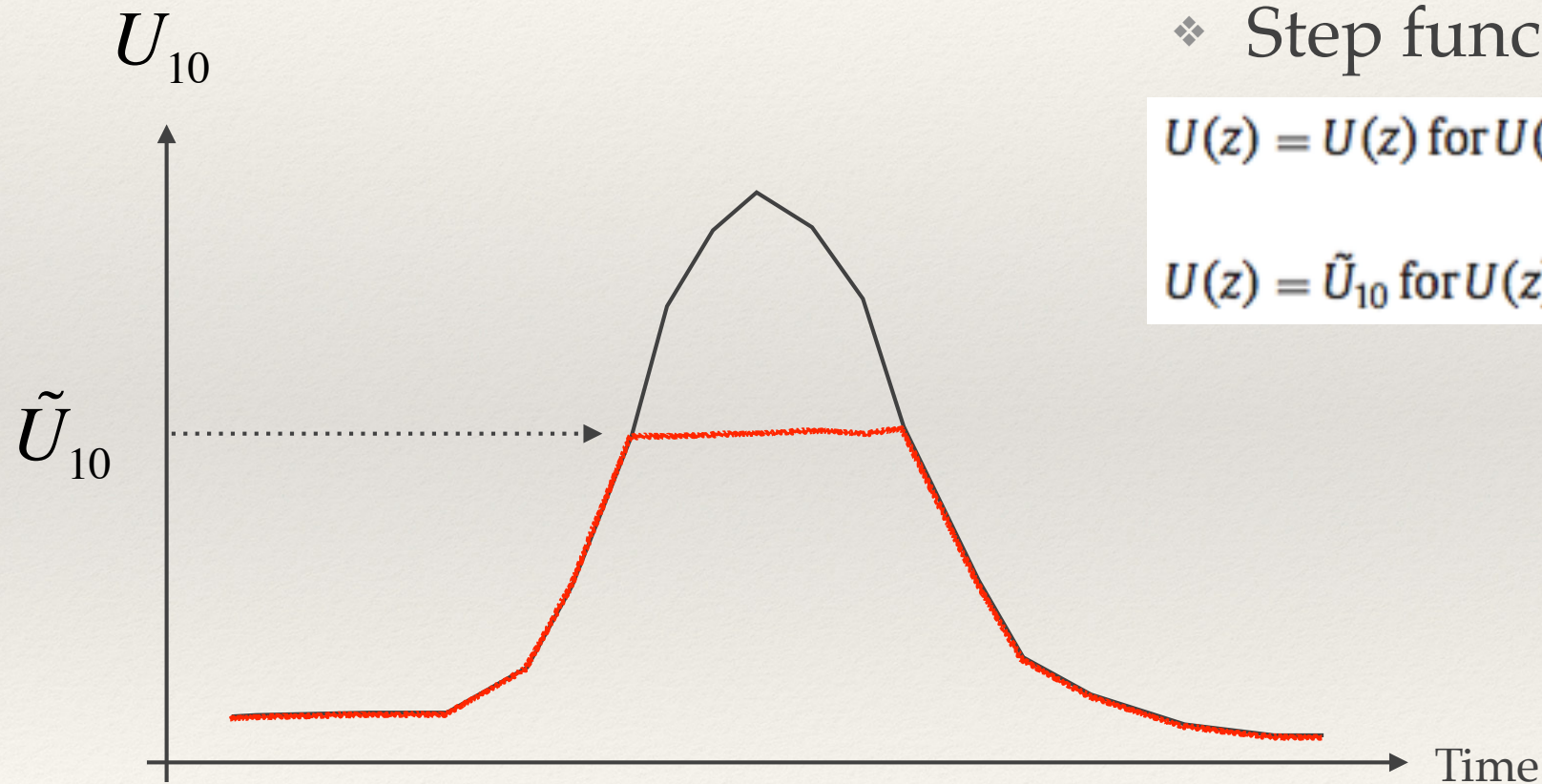
Kim et al., 2015 *Ocean modelling*

- ❖ Threshold for a levelling off based on measurements
 - ❖ 33 *m/s* : Powell et al. (2003)
 - ❖ 30-40 *m/s* : Donelan et al. (2004)
 - ❖ 22-23 *m/s* : Black et al. (2007)



Levelling Wave dependent Cd off in the exponential term due to Typhoon Hainan 2013

- ❖ Wind profiles only in the exponential term



Levelling off in the exponential term

- ❖ Levelling off the wave dependent Cd in the exponential wave growth term

$$U(z) = U(z) \text{ for } U(z) < \tilde{U}_{10}$$

- ❖ Step functions

$$U(z) = \tilde{U}_{10} \text{ for } U(z) \geq \tilde{U}_{10}$$

- ❖ Wind profile

$$U(z) = \frac{u_*}{k} \ln \left(\frac{z + z_e + z_0}{z_e} \right), \text{ if } U(z) < \tilde{U}(z)$$

$$\tilde{U}(z) = \frac{u_*}{k} \ln \left(\frac{z + z_e + z_0}{z_e} \right), \text{ if } U(z) \geq \tilde{U}(z)$$

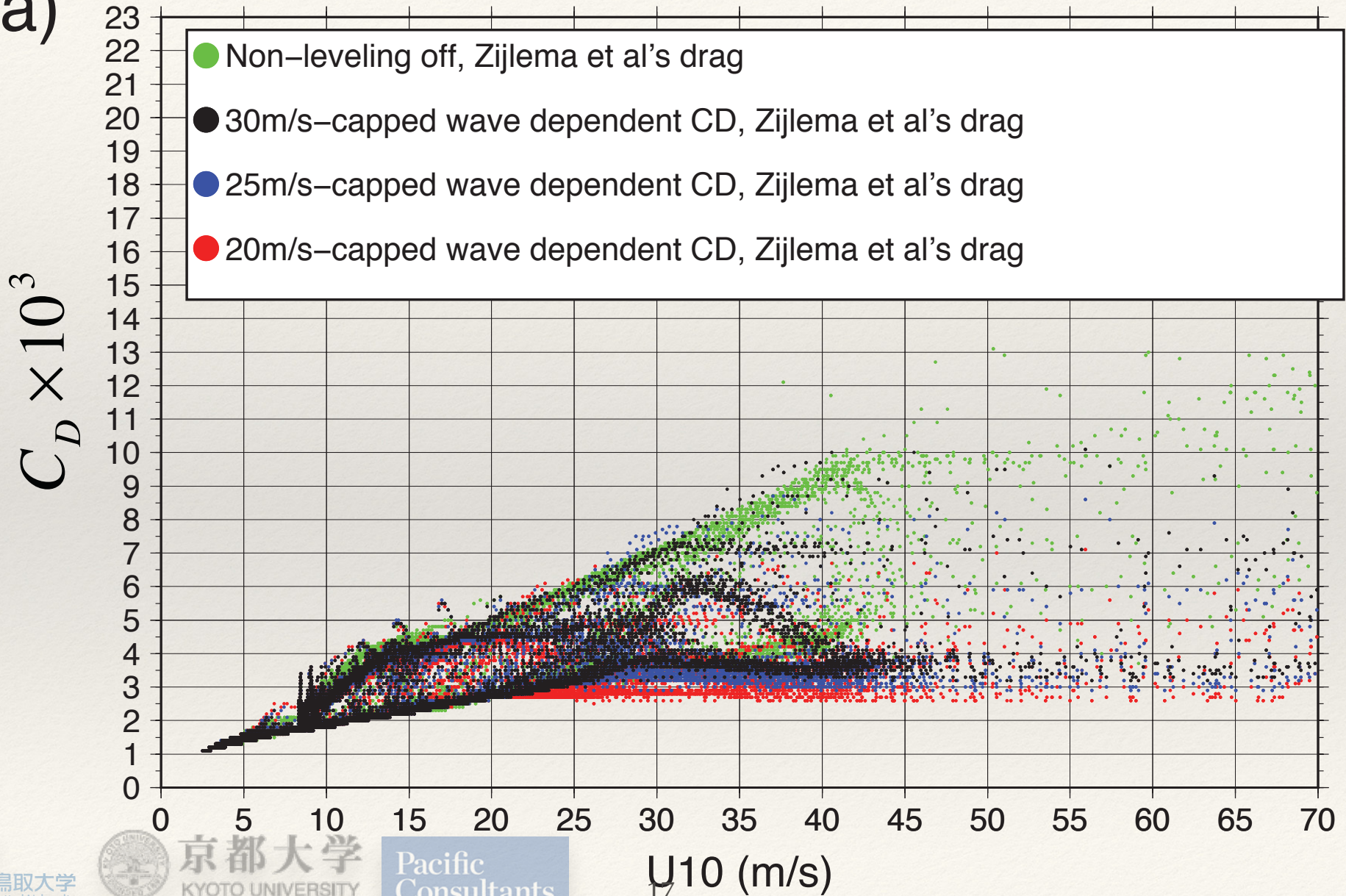
- ❖ Wind speed-capped
Wave dependent Cd

$$C_d = u_*^2 / U(z)^2 = \kappa / \ln \left(\frac{z + z_e + z_0}{z_e} \right), \text{ if } U(z) < \tilde{U}(z)$$

$$C_d = u_*^2 / \tilde{U}(z)^2 = \kappa / \ln \left(\frac{z + z_e + z_0}{z_e} \right), \text{ if } U(z) \geq \tilde{U}(z)$$

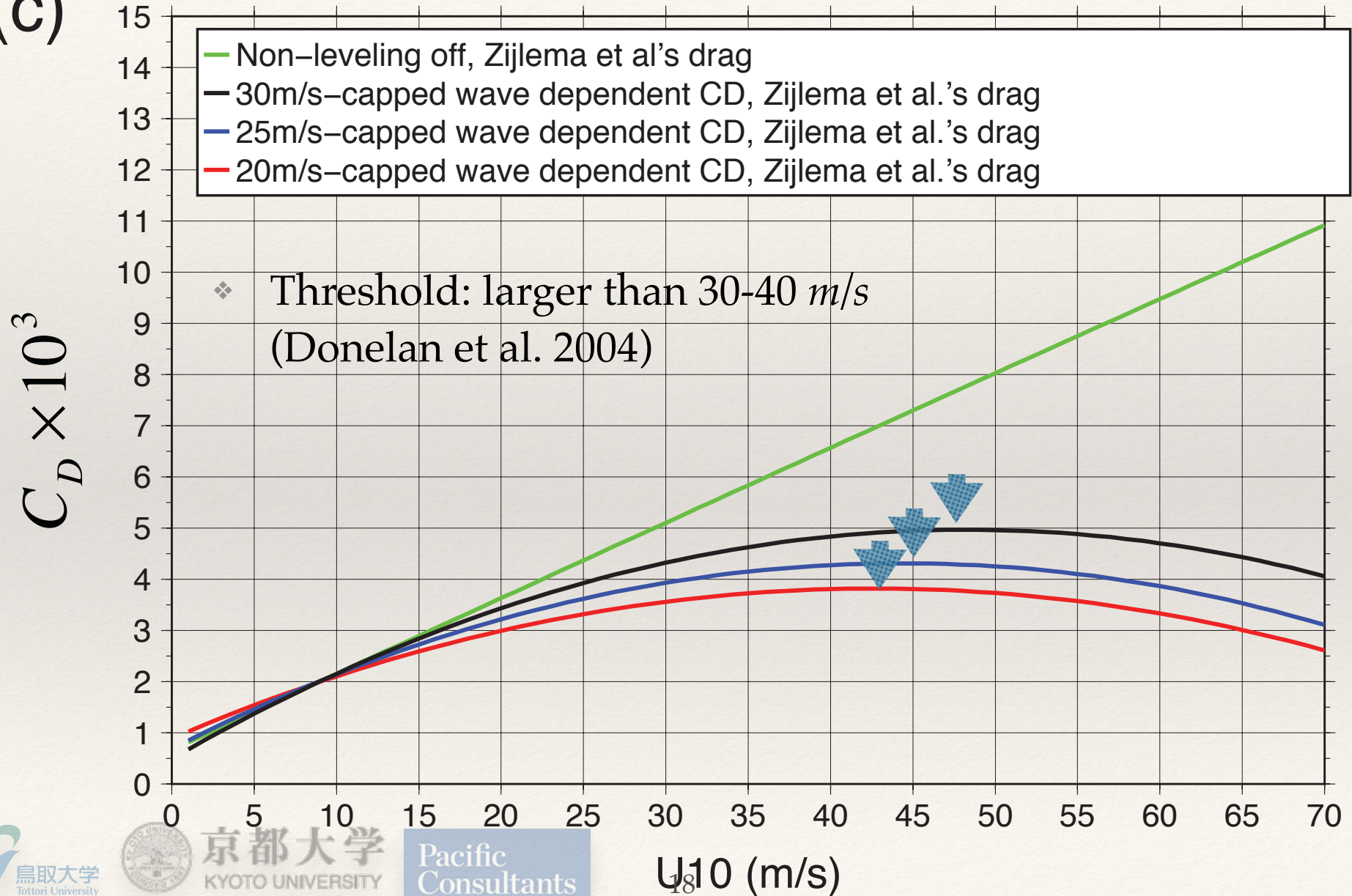
Scattered wave dependent C_d with levelling off due to Typhoon Hainan 2013

(a)



The best-fitted wave dependent C_d to the 2nd-order polynomial due to Typhoon Hainan 2013


(C)



The wave¤t interaction-induced bottom drag, f_c

- ❖ Conventional method

- ❖ Manning number, n ,

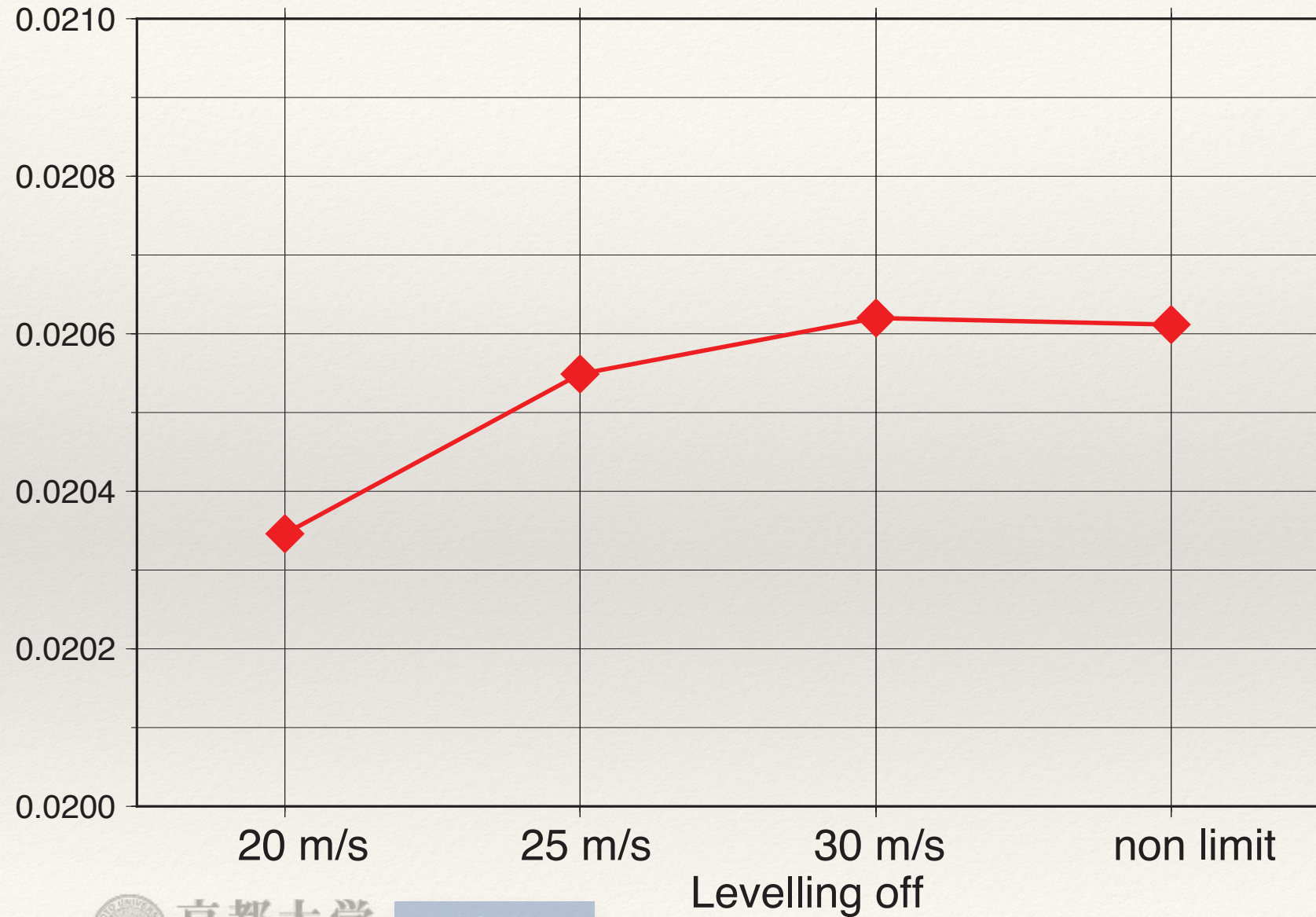
$$\tau_b = \rho_w g n^2 \frac{|\vec{Q}|}{h^{7/3}} \quad f_c = 8 \times \frac{g n^2}{h^{1/3}} \quad \tau_b = \rho_w \frac{f_c}{8} \frac{|\vec{Q}|}{h^2}$$


- ❖ *Signell et al., 1990 & Davies and Lawrence, 1995*

$$k_{bc} = k_b \left[C_1 \frac{U_{*cw} A_b}{U_w k_b} \right]^\beta \quad f_c = 2 \left[\frac{K}{\ln(30 z_r / k_{bc})} \right]^2$$



Averaged Manning Number converted from f_c due to Typhoon Haiyan 2013

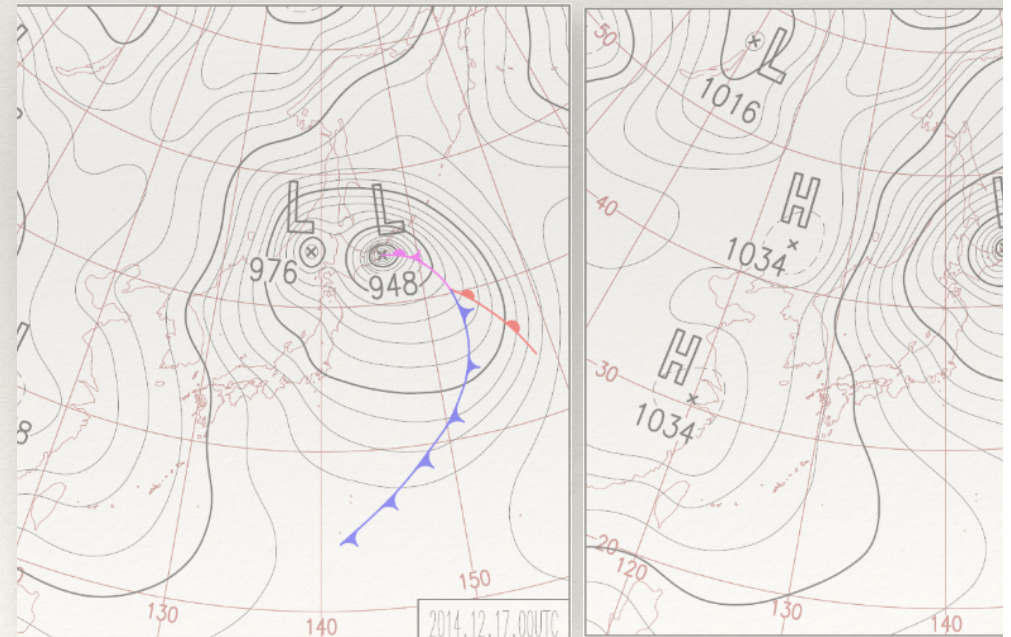
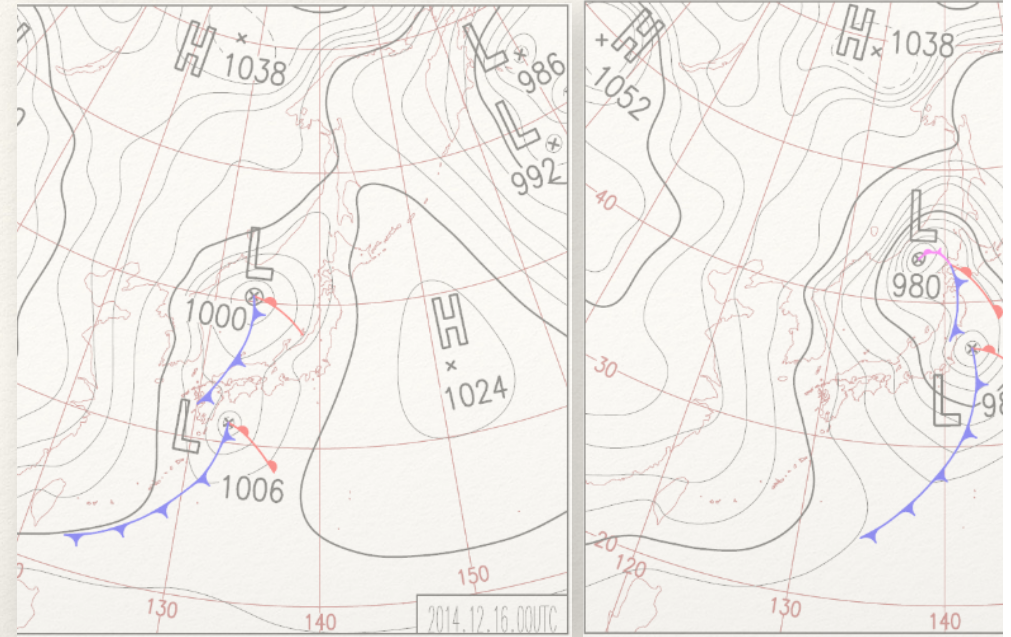




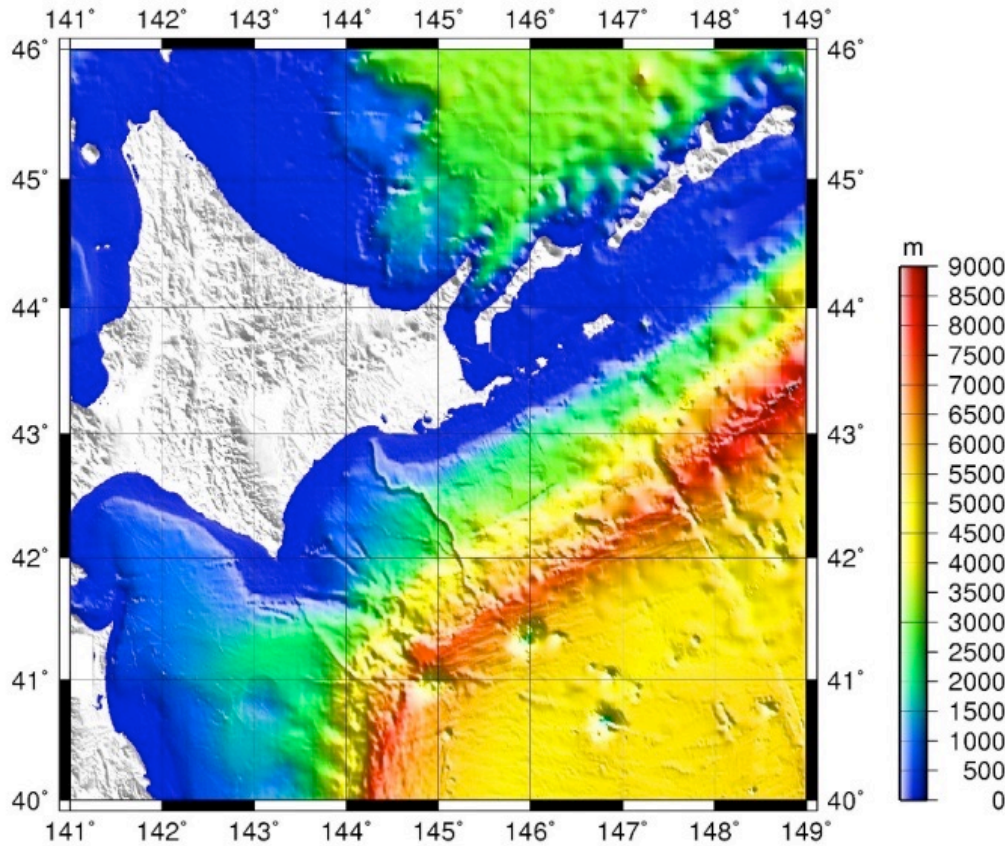
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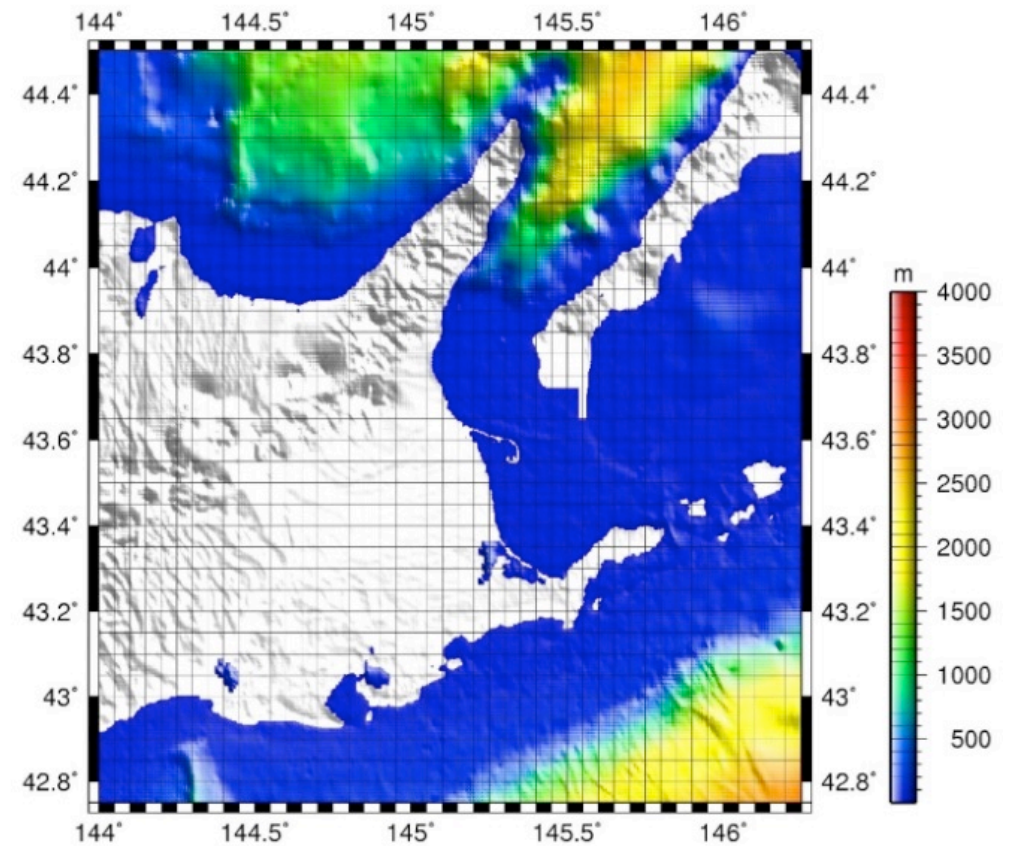
Cal. Conditions



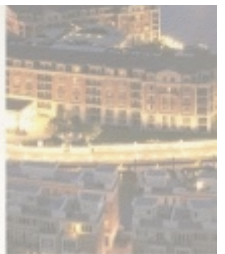
Calculation domains



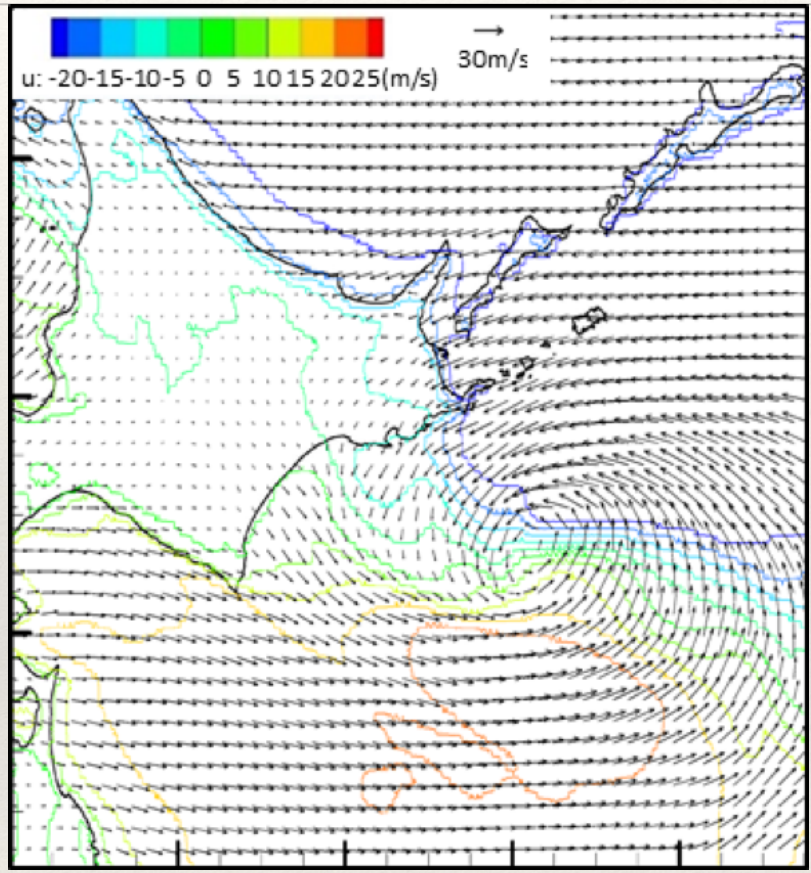
Grid size: 677 m x 927 m



Grid size: 226 m x 309 m

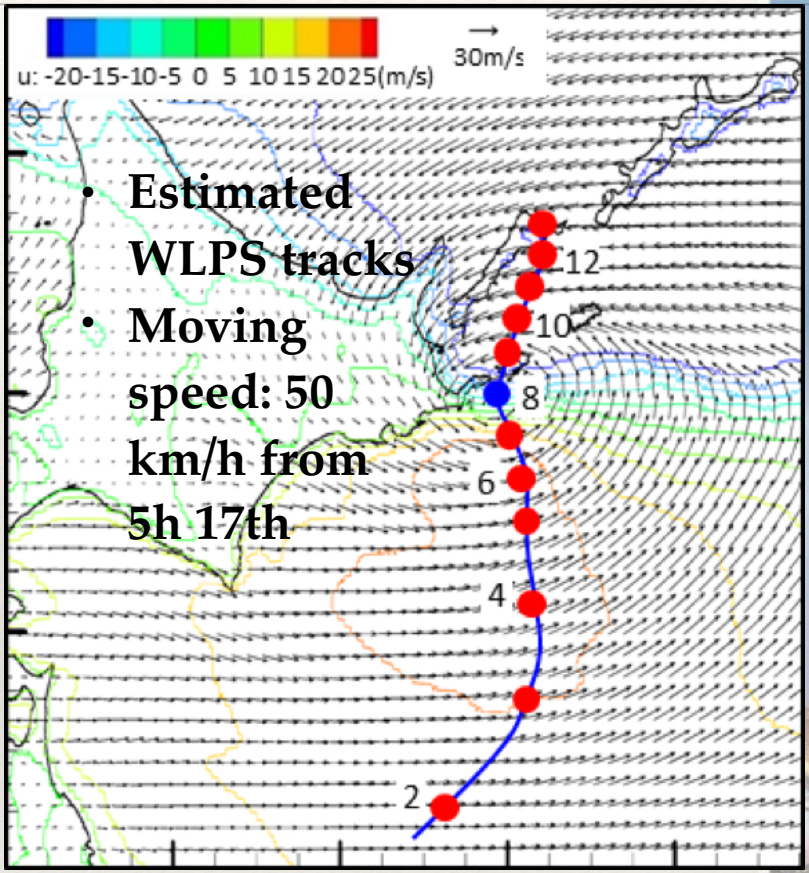


Wind and pressure fields by JMA



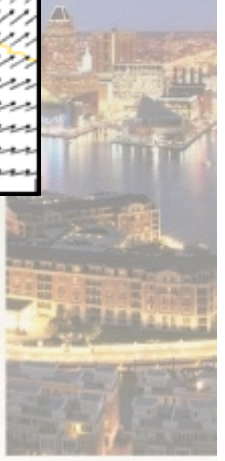
5:00 17th Dec.

- Max. Wind speed: 30.7 m/s



8:00 17th Dec.

- The wind and pressure field
 - when the minimum drop of pressure is observed in Nemuro.

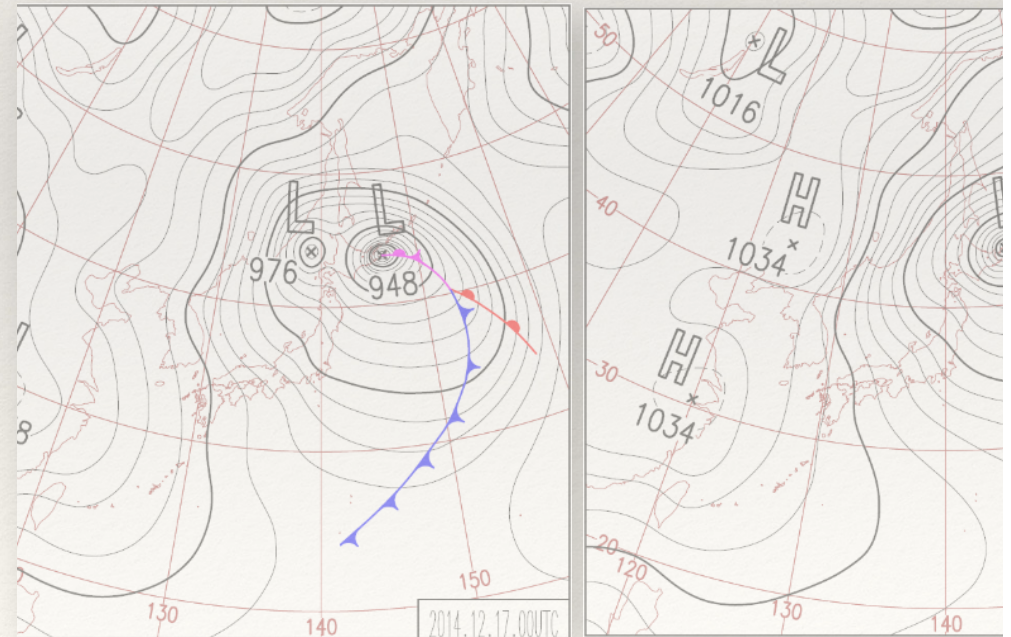
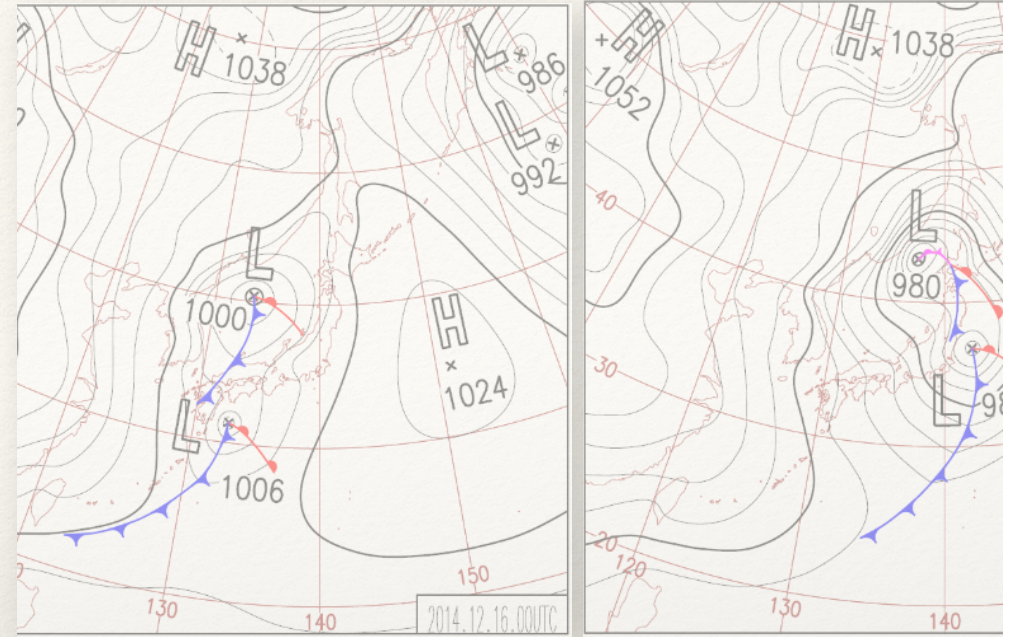


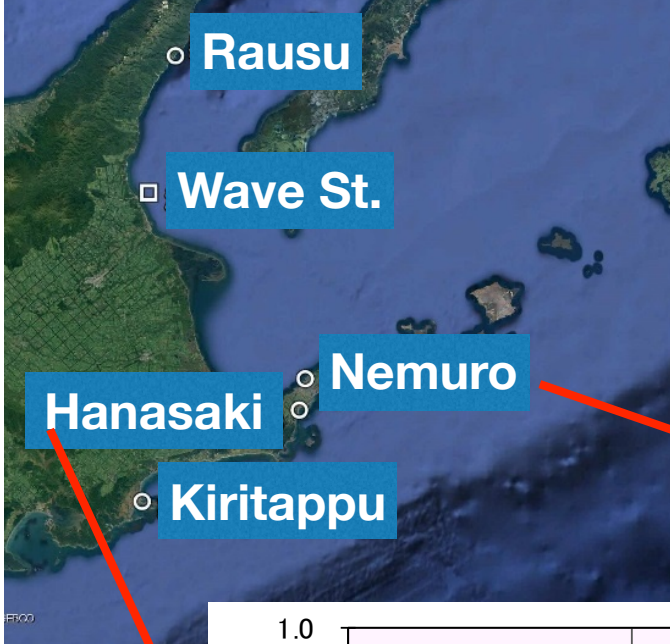


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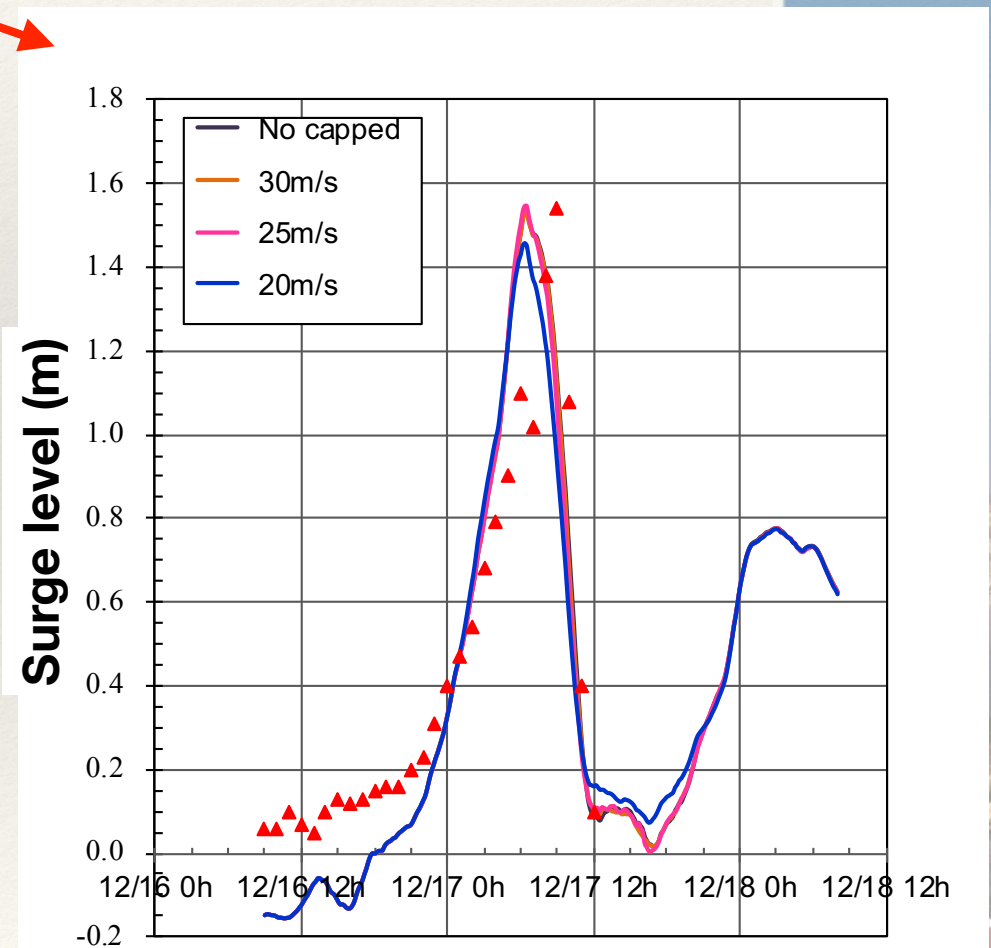
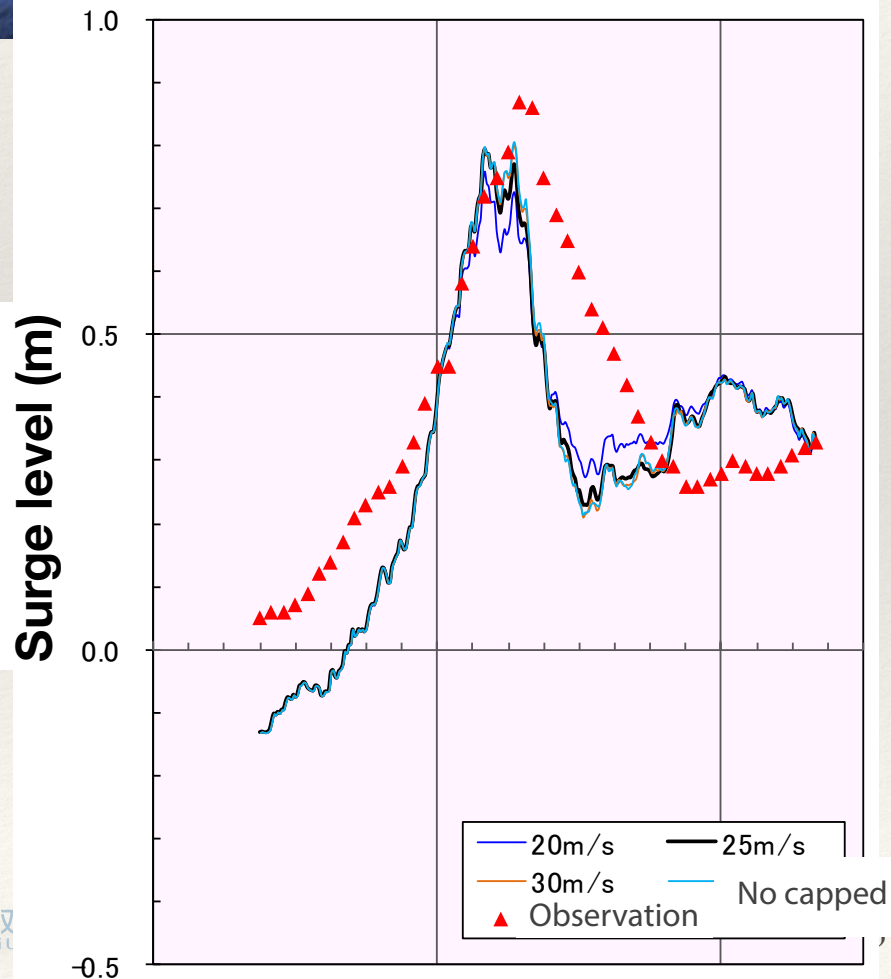
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Results

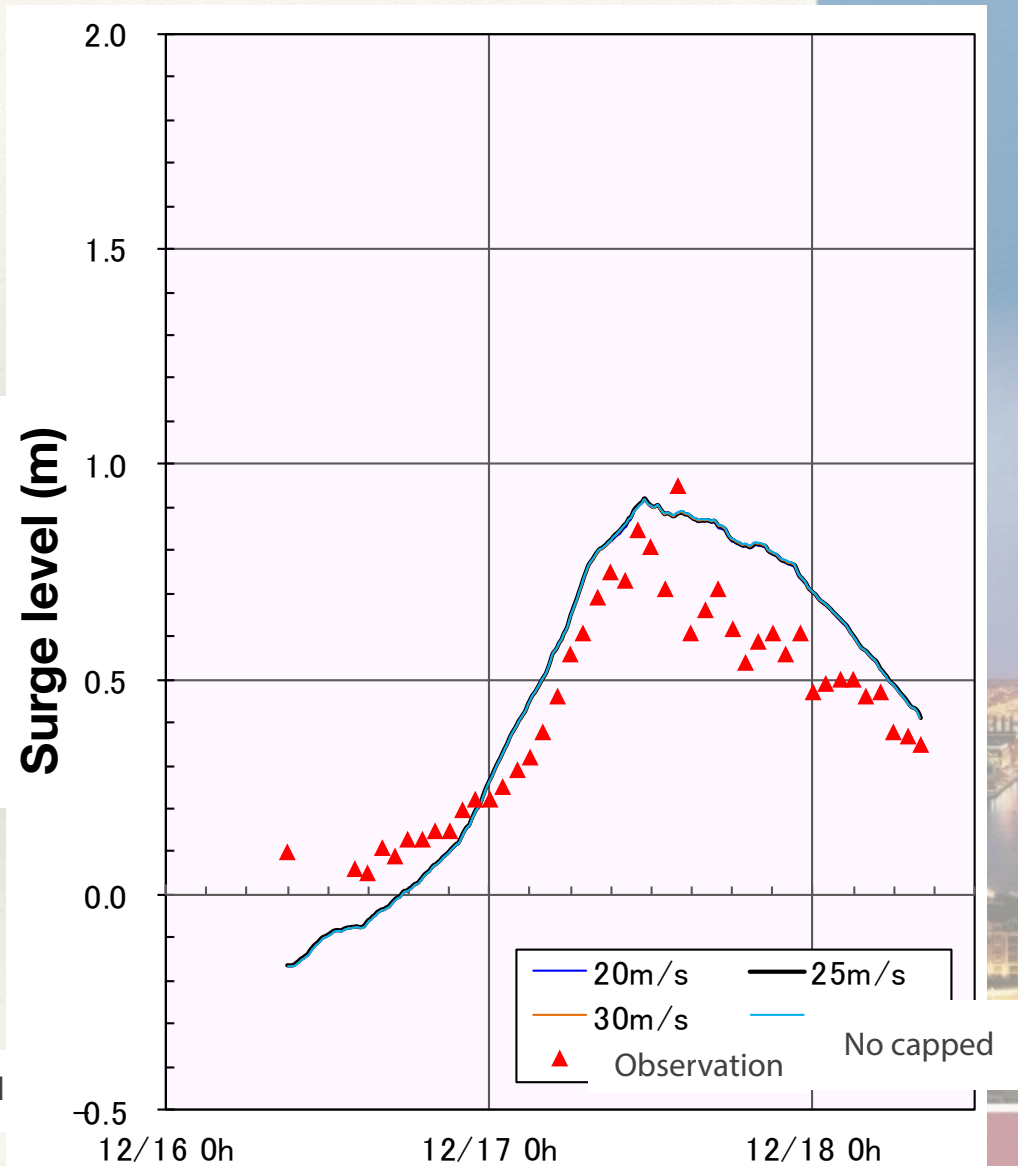
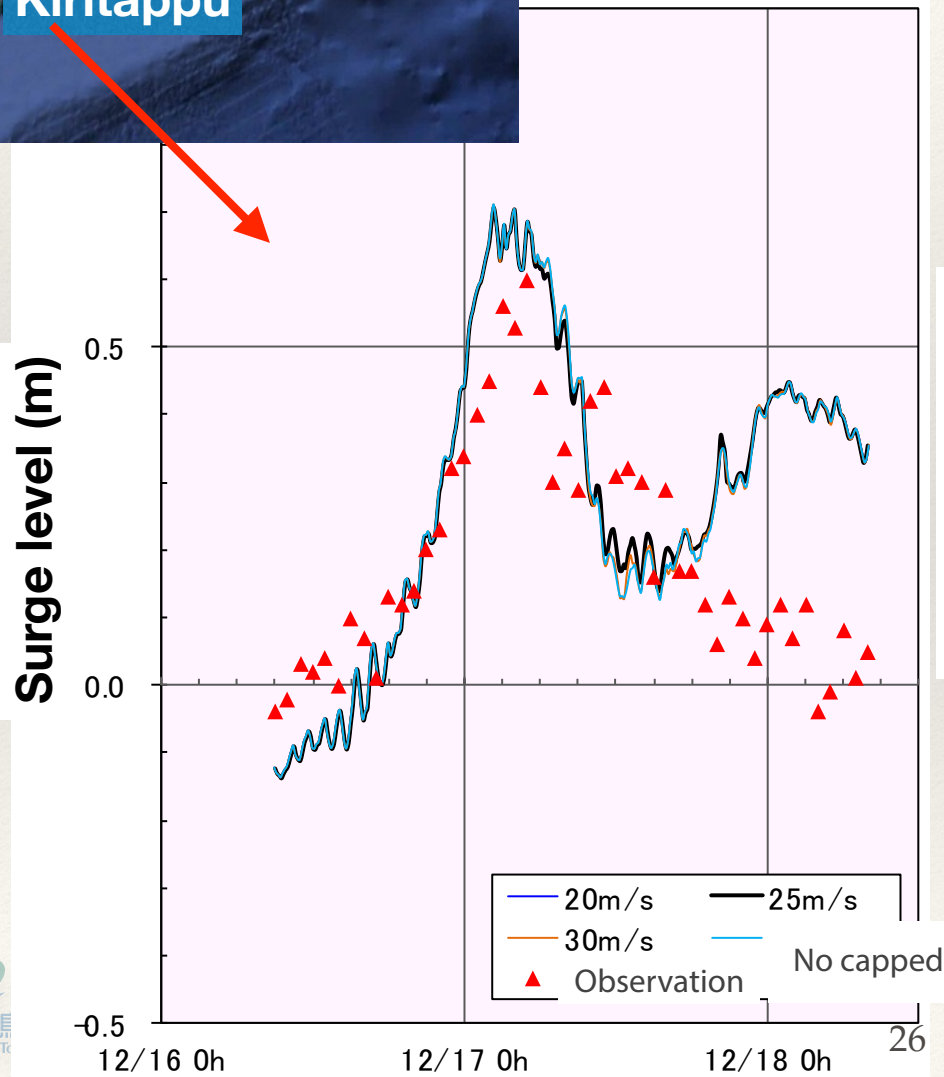
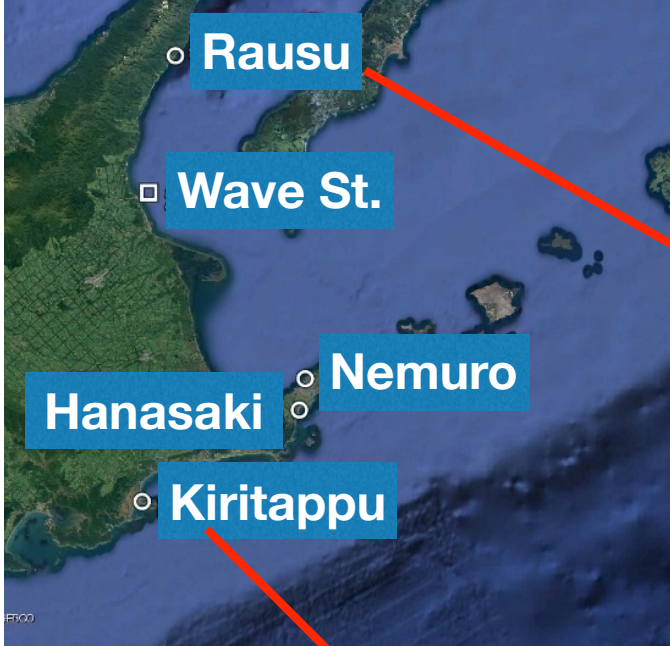


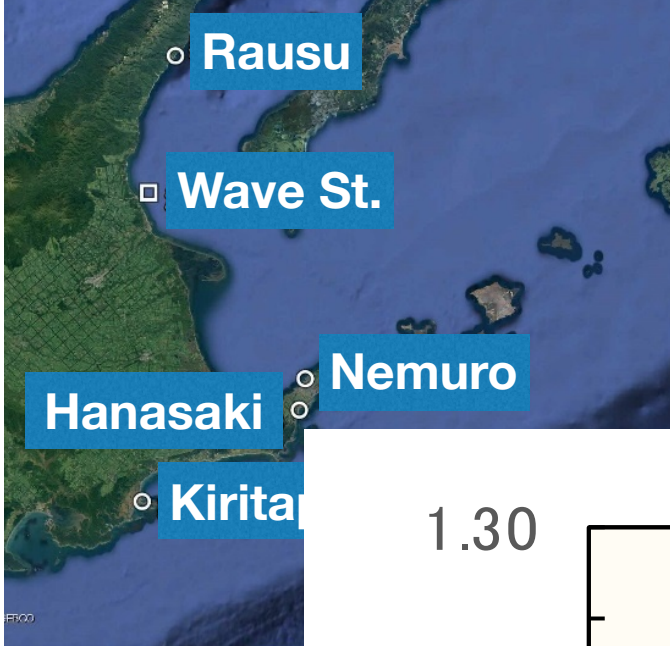


Results: Surge levels



Results: Surge levels

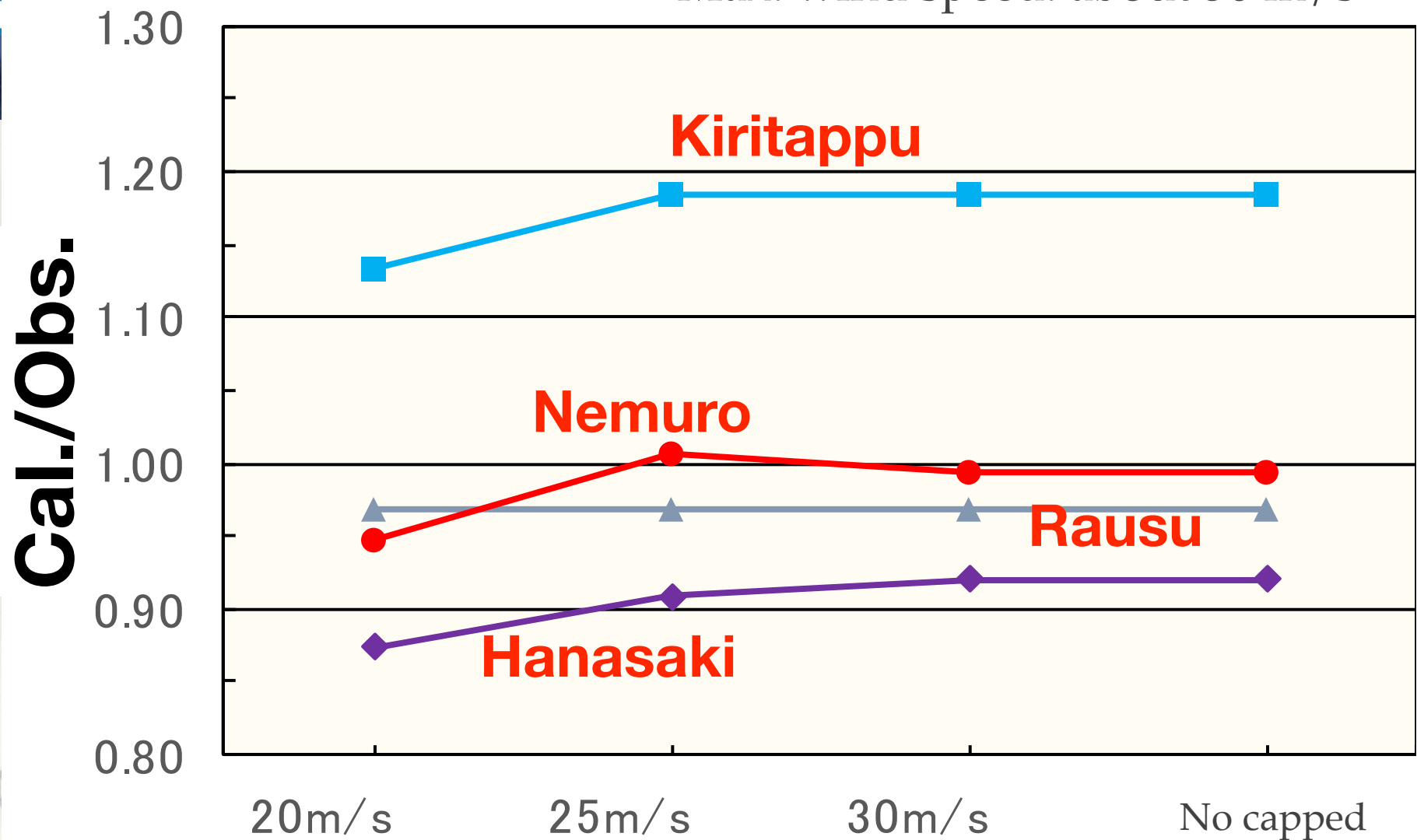




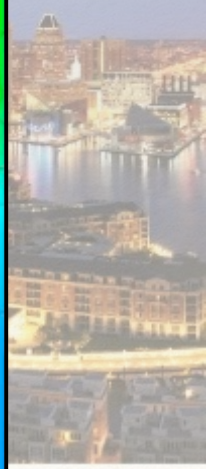
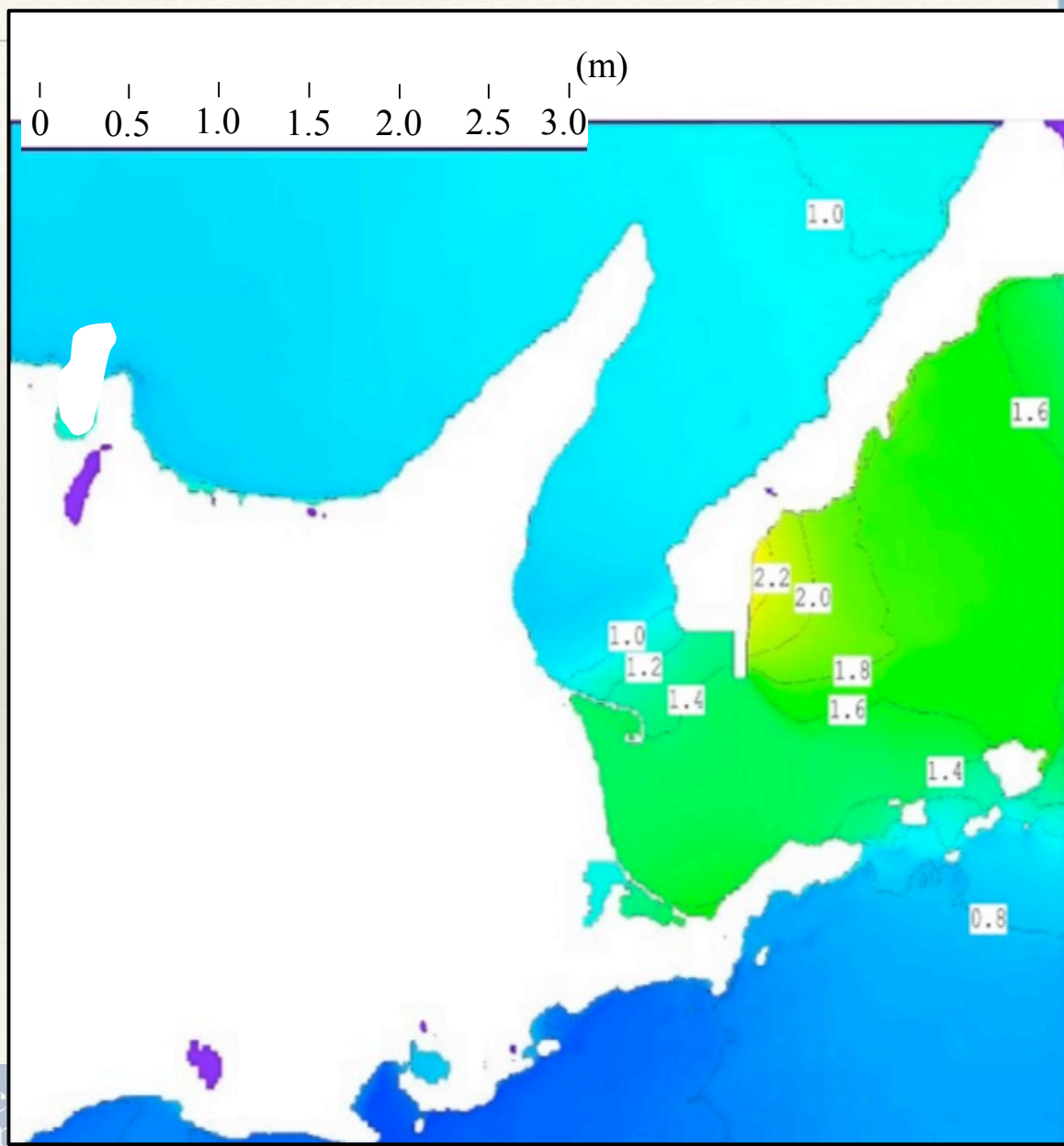
Ratio of cal. / Obs. for Max. level



Max. wind speed: about 30 m/s

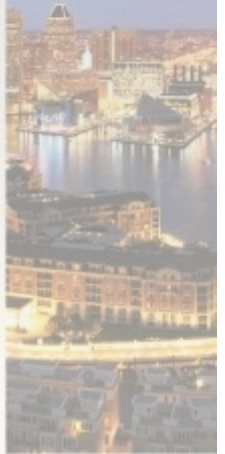
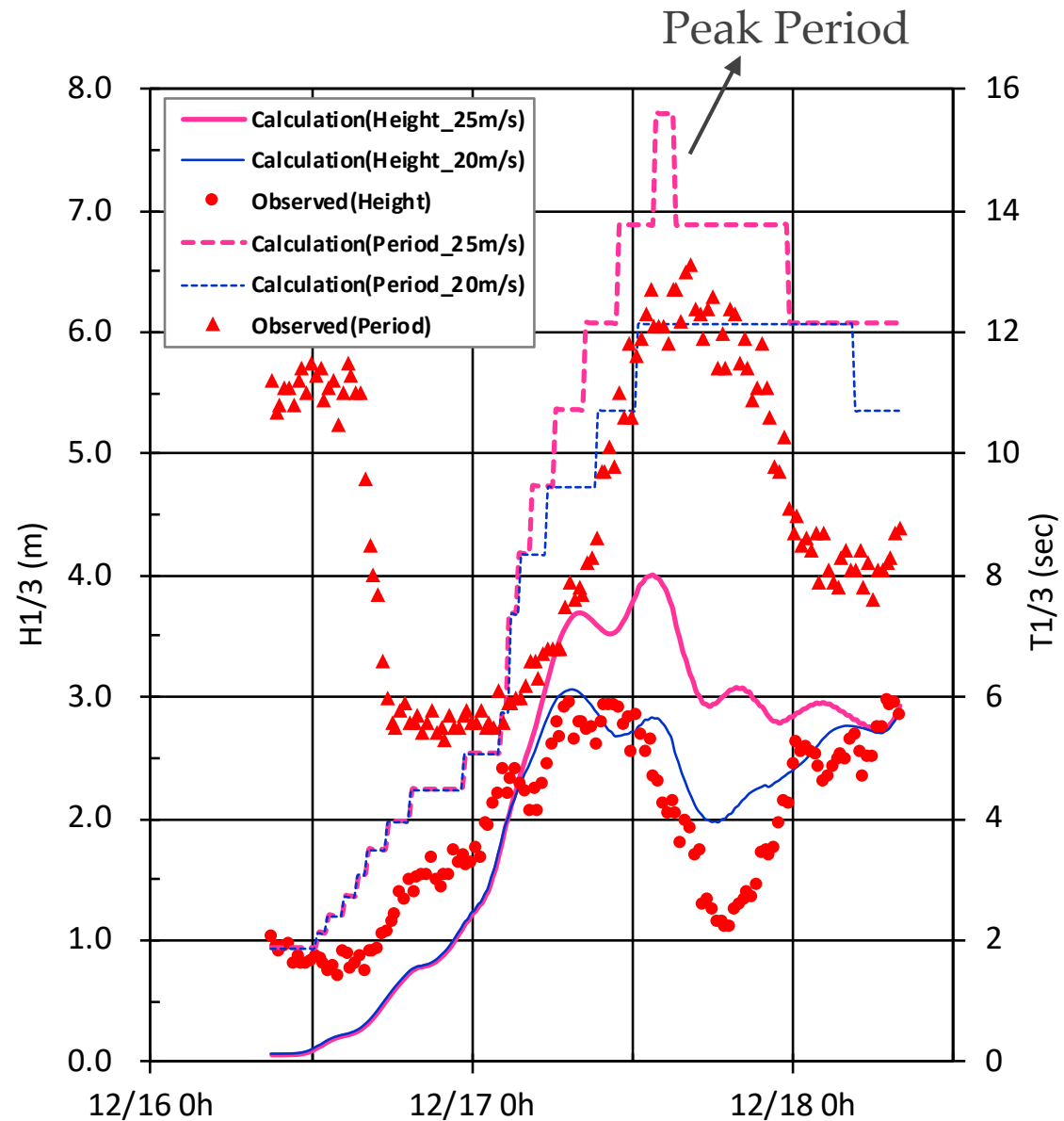


Results: Max. surge levels (25 m/s capped)



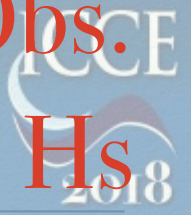


Results: Waves



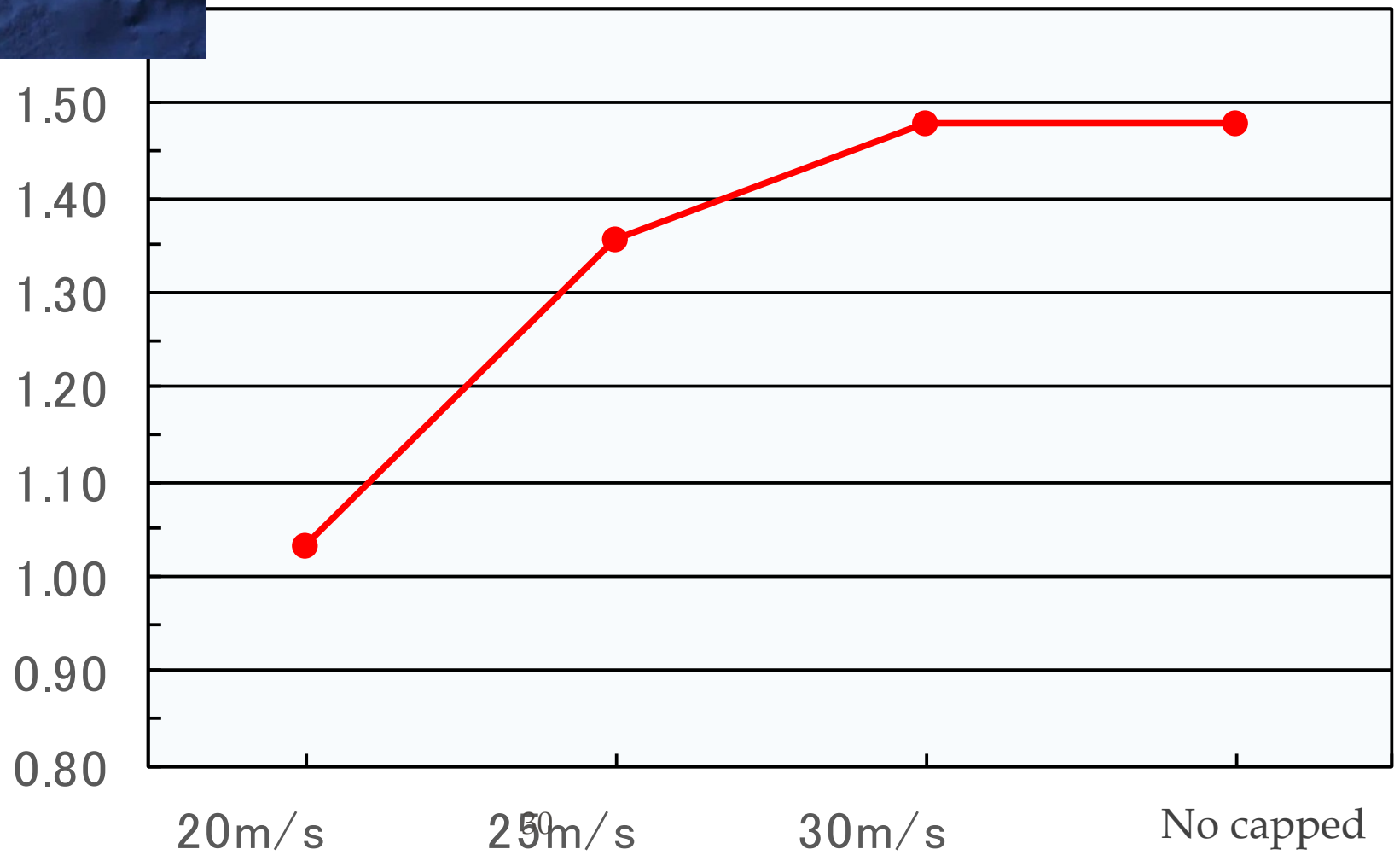


Ratio of cal. / Obs. for Max. Hs

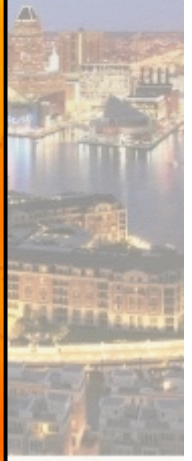
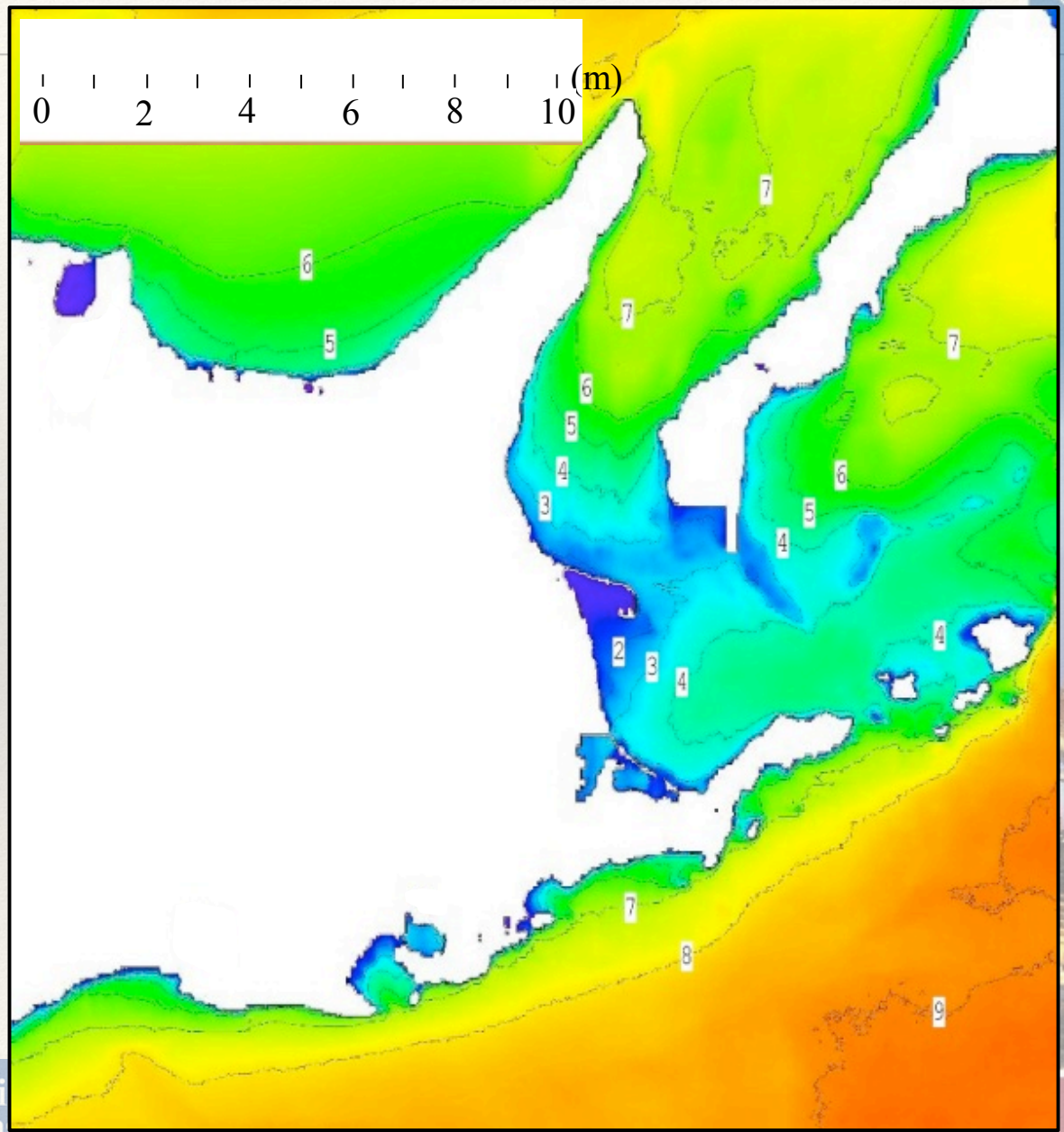


Max. wind speed: about 30 m/s

Cal./Obs.



Results: Max. Hs (20 m/s capped)



Summaries

- ❖ Winter Low Pressure System (WLPS) 2014 induced the record-break surge level in Hokkaido.
- ❖ A series of surge and wave coupled simulations was conducted by using the coupled model of SuWAT
 - ❖ wind speed capped wave dependent drag coef.
 - ❖ wave-current interaction-induced bottom drag coef.
- ❖ Leveling off at 20 m/s was best for waves at one station
- ❖ Leveling off at 25 m/s was best for surges at 4 station
- ❖ Discrepancy of the specific wind speed for the leveling off has to be investigated.
 - ❖ tides?
 - ❖ more observed wave data?
 - ❖ stronger typhoons/hurricanes?



Questions or comments ?

Thank you very much

