

# **HIGHLY RESOLVED DIRECTIONAL PROPERTIES OF WIND WAVES AND SWELL WITH VARIOUS SCALES**

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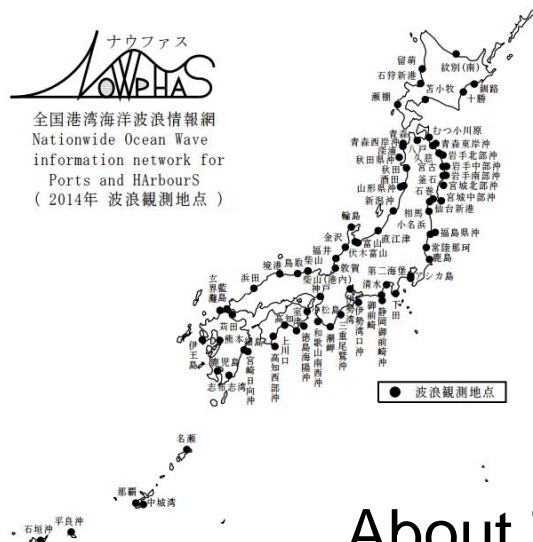
- Aim of wave observation and its applications in Japan
- Estimation of directional spectrum with Bayesian inversion
- Results:
  - Agreement with  $3/2$  power of law for wind waves
  - Relationship between Non-dimensional wave statics and directional spreading
- Discussions and conclusions
- Future works

# Wave Observation by MLIT

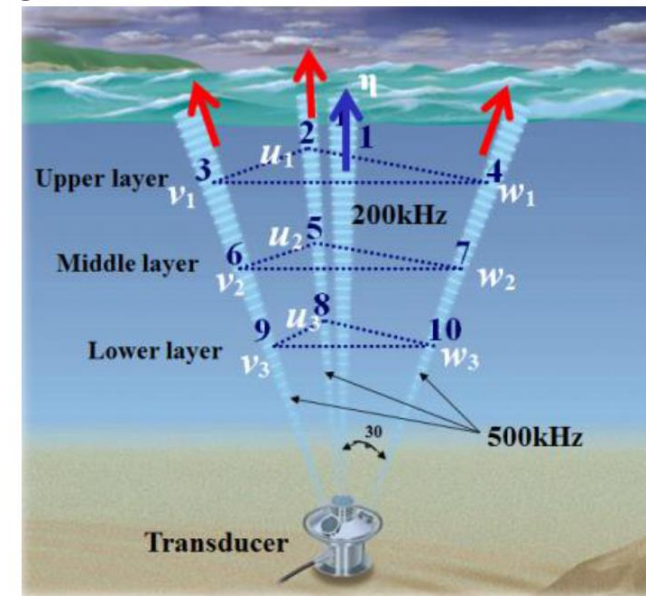
- Observation plan designed by MLIT  
(**M**inistry of **L**and, **I**nfrastructure and **T**ourism)
- Started with pressure gauges in 70's
- Two major type instruments at present  
Buoy type: GPS sensor on mooring buoy  
bottom installed type: Wave-ADCP, USW, pressure gauge



GPS-buoy



About 70 site



Wave-ADCP

# Aim of Wave Observation and Its Applications in Japan

Wave statics (height, period, direction) in various scales

are used for coastal design, coastal operation, etc.

- **High sea** (wave height > 8m) -> Sampling of reference data for design use
- **Moderate sea** ( wave height < 1m) -> Monitoring of tranquility for marine activities
- **All time** (from low to high) -> validation of model prediction using wave forcing, e.g., shoreline change, wave climate, etc.

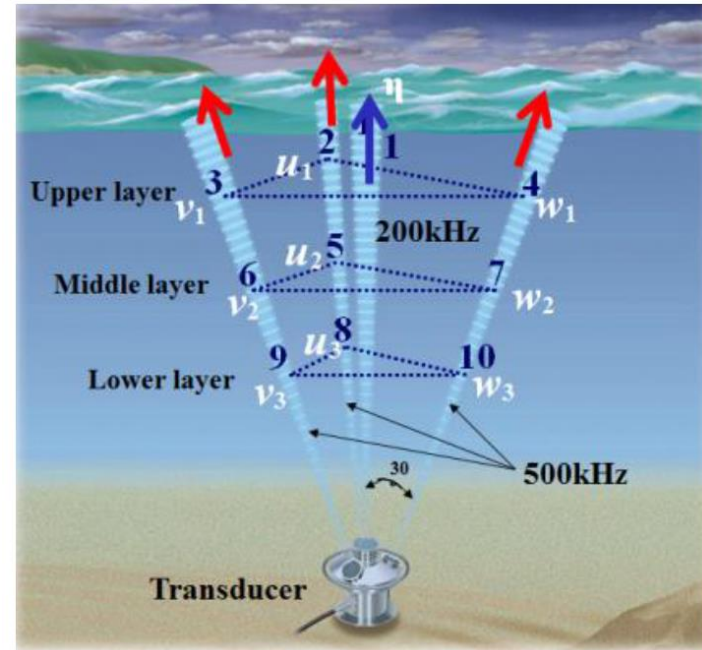
-> These aims requests instruments for moderate sea to high sea

# Observation from Moderate Sea to High Sea

- Bottom installed type is desirable for its high durability to high sea state  
    <- wave staff on sea surface can be troubled easily by high waves
- Array instruments is required for directional information  
    <- directional info is obtained through directional spectrum

Wave-ADCP meets these requests

- function :
  - vertical ultra sonic wave -> surface displacement
  - along beam back scattering -> current profile
- Array observation with single instrument



Wave-ADCP

# Inversion of Directional Spectrum

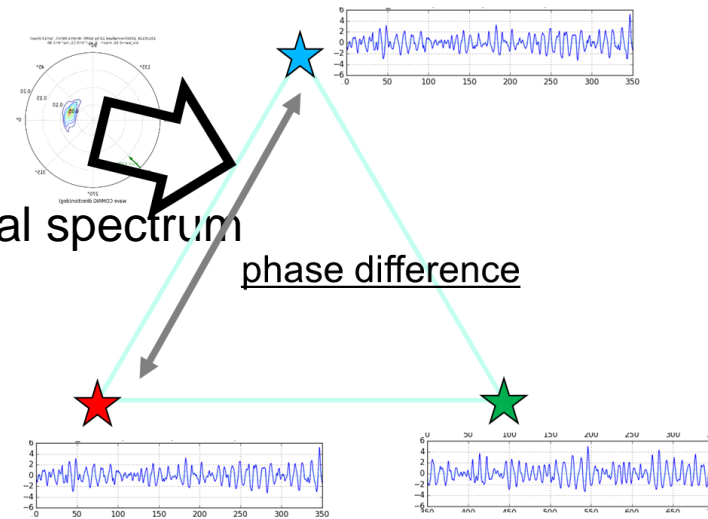
- Discretized Integral equation for inversion of directional spectrum

$$d = G \cdot m$$

$d$ : observational data vector(phase difference)

$G$ : model matrix(transfer function)

$m$ : directional spectrum



Simple solution by Least Square Method:

$$m^* = G^{-1} \cdot d$$

Problems in simple inversion method

- Low resolution due to  $G$ , which is transfer function of Doppler velocity
- Observational noise in  $d$  make estimator unstable
- Increasing size of  $d$  give Collinearity

} Ill-Posed problem

# Bayesian Inversion of Directional Spectrum

- Bayesian Inversion enables smoothing based on prior belief
- Introduction of smoothing by Bayesian inference (Hashimoto, 1987)
  - smoothing by minimizing second difference on directional spectrum

Hyper parameter  $u$ , which controls contribution of prior distribution, is determined by ABIC(Akaike Bayesian Information Criteria) automatically

$$R(\mathbf{m}) = |\mathbf{d} - \mathbf{G} \cdot \mathbf{m}|^2 + \underline{u^2 |\mathbf{L} \cdot \mathbf{m}|^2} > \min$$

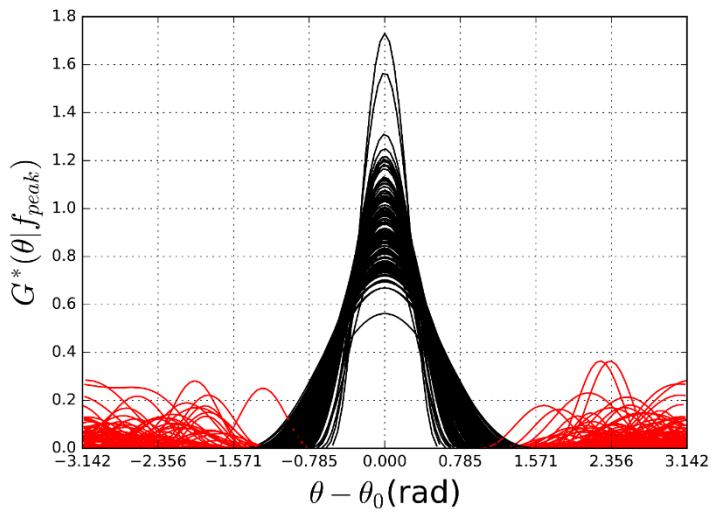
penalty term for smoothing(varying with  $u$ )

$u \rightarrow 0$ : solution by Least Square Method

$u \rightarrow \infty$ : uniform directional spectrum(NO contribution of observed data)

# Partitioning and Identification of Several Wave Systems

Major peak + Minor peaks

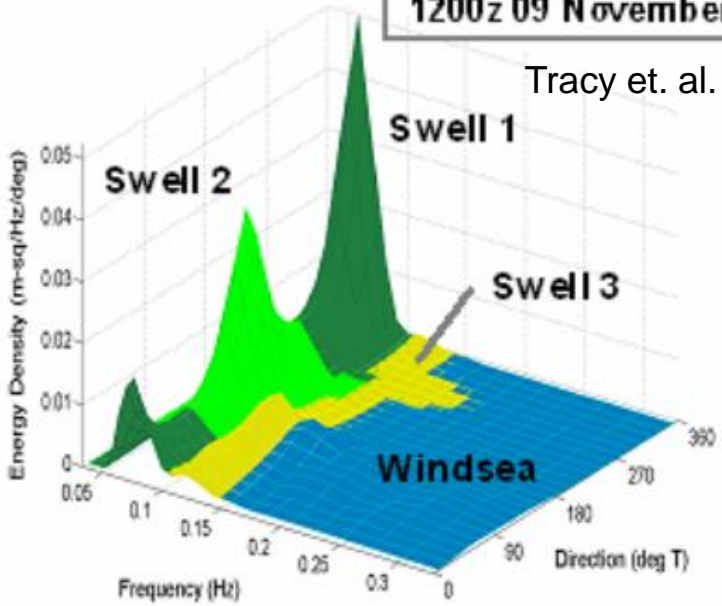


- Observed directional spectrum can show several peaks
  - Detailed analysis for each peaks is required
- WITS (Wave Identification and Tracking System) by Hanson et al. 2001

- peak detection by **Water Shedding Algorithm**
- merging several peaks close to each other
- rejecting apparent peaks with small energy

**Station 51028**  
1200z 09 November

Tracy et. al. 2007



- calculate wave statics for each peaks
  - height, period, direction
- Identification for windsea and swell by inverse wave age,  $|U_{10} \cos \delta|/C_p$

$U_{10}$ : 10m-above local wind(m/s)

$\delta$ : Directional difference between wave and wave(rad)

$C_p$ : Phase speed at spectral peak(m/s)



# Data Processing

- site:  
Akita (Japan sea, **windsea** is dominant in winter)  
Hachinohe (Pacific ocean, **swell** is dominant in summer)
- analysis method: directional spectrum by Bayesian inversion
- criteria to reject marginal data for detailed analysis :

$$H_{m0}(\text{m}) : 0.5 \sim 6.0$$

$$T_p(\text{s}) : 4.0 \sim$$

reject counter swell

$$|D_{wind} - D_{wave}| (\text{deg}) : < 45.$$

reject transient wave

$$\Delta U_{10} (\text{m/s}) : < 5.0$$

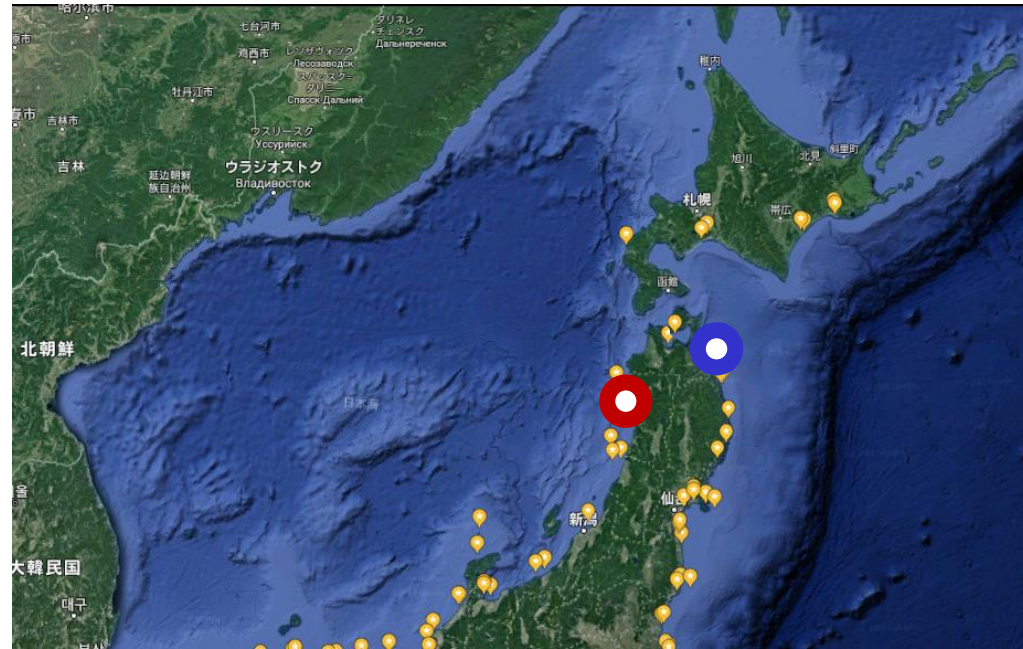
$$\Delta D_{wave} (\text{deg}) : < 30.$$

$$\Delta D_{wind} (\text{deg}) : < 30.$$

separating windsea and swell

$$\text{Windsea: } U_{10}/C_p(-) : \mathbf{1.0} \sim$$

$$\text{Swell: } U_{10}/C_p(-) : \mathbf{0.1} \sim \mathbf{0.5}$$



# Summary of samples

wind sea(50/26280)

$H_{m0}$ (m) : 0.62 ~ 5.21

$T_p$ (s) : 4.4 ~ 11.5

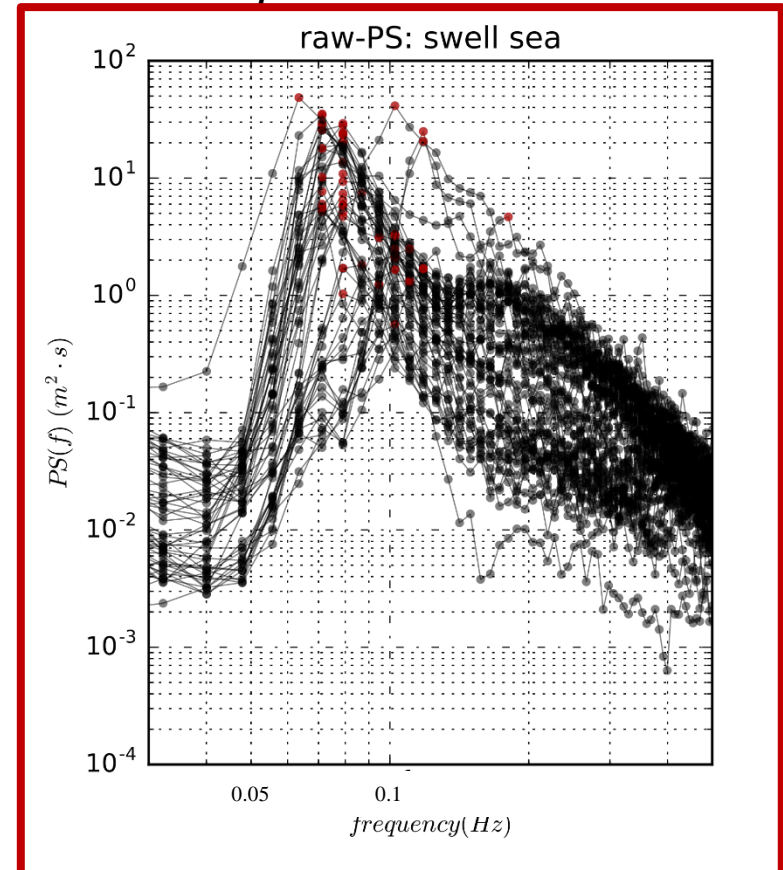
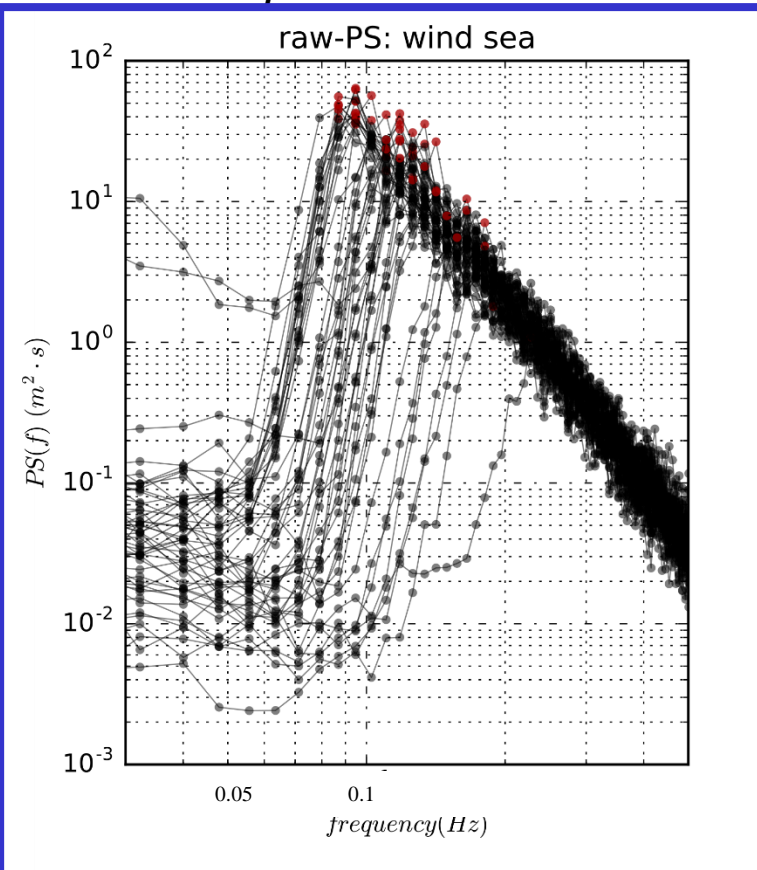
$U_{10}/C_p$ (-): 0.93 ~ 1.69

swell sea(53/26280)

$H_{m0}$ (m) : 0.50 ~ 4.18

$T_p$ (s) : 5.8 ~ 15.8

$U_{10}/C_p$ (-): 0.16 ~ 0.50



# Results: validation on Power Law

- The 3/2 power of law for wind wave
  - Non-dimensional wave height and period

$$H^* \sim T^{*3/2}$$

$$H^* = gH_{m0}/|U_{10} \cos \delta|^2$$

$$T^* = gT_p/|U_{10} \cos \delta|$$

$U_{10}$ : 10m-above local wind(m/s)

$\delta$ : Directional difference between wave and wave(rad)

- Validation with samples with various scales, including moderate sea to high sea
- Does swell also conform the 3/2 power of law ?

Suzuki(2011)

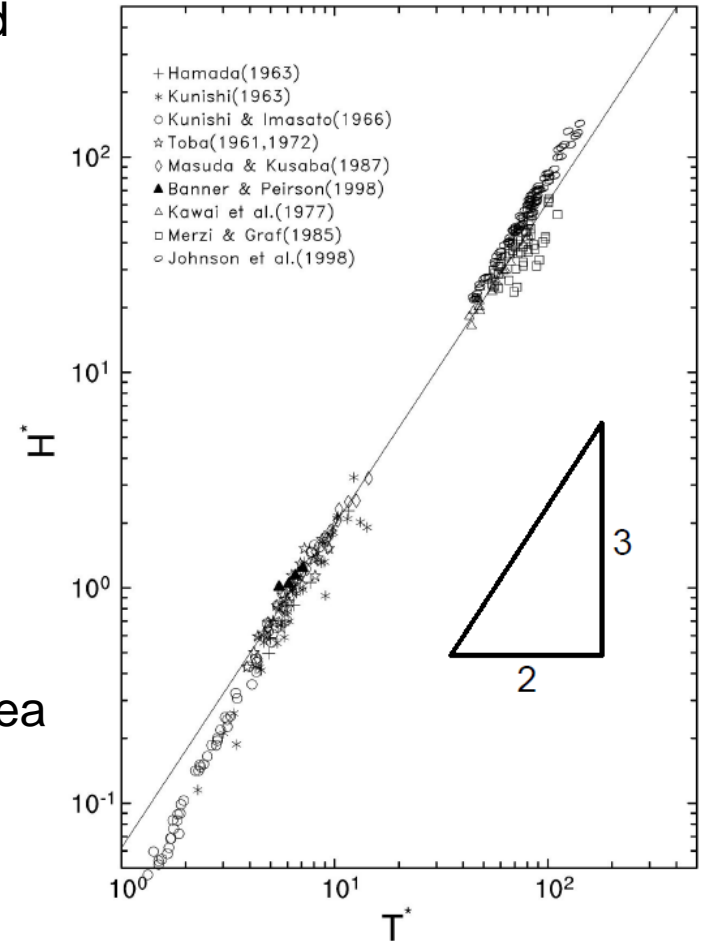
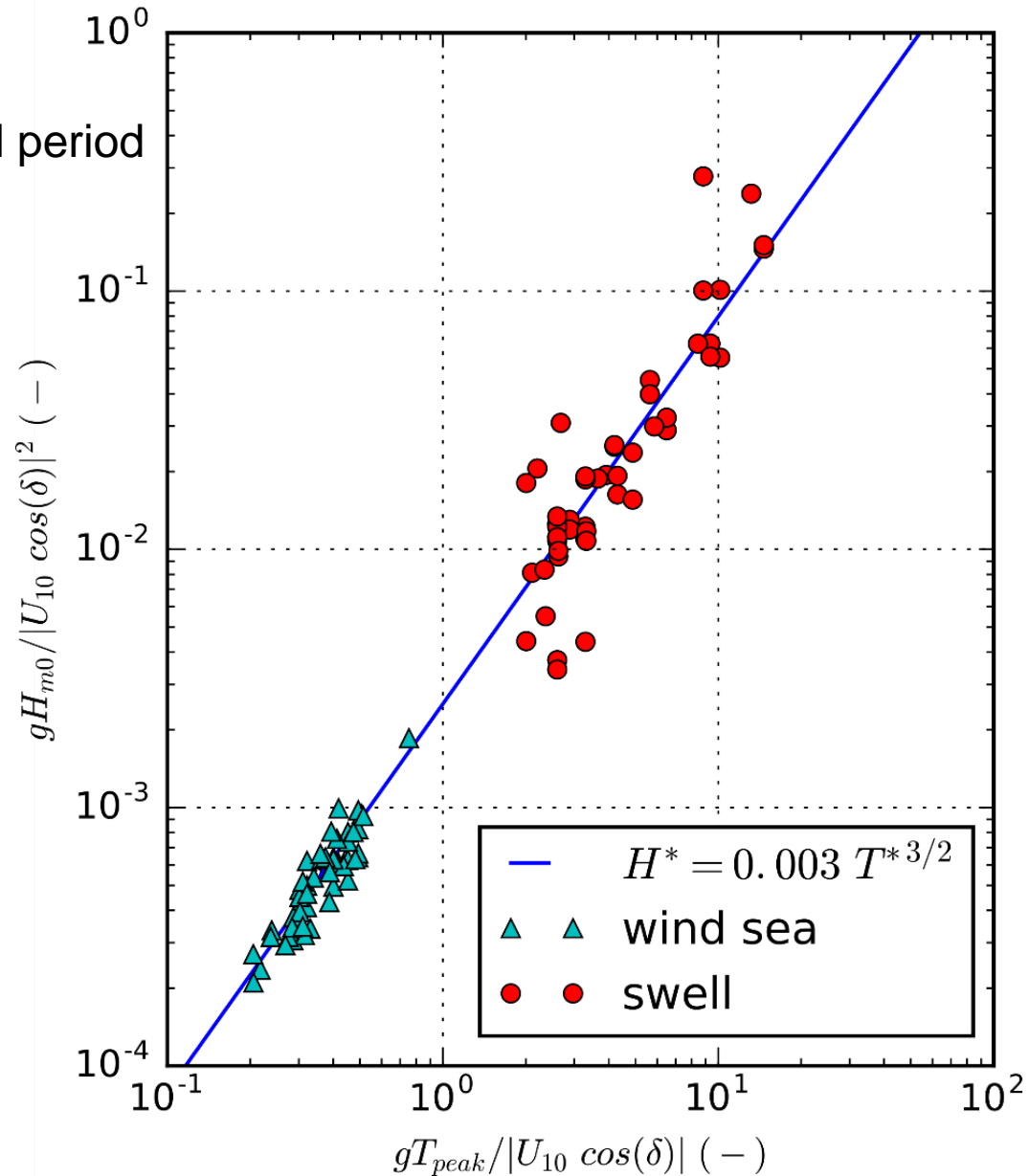


Fig.5 The relation between non dimensional wave height  $H^*$  and period  $T^*$ . The 3/2 power law (Toba, 1972) is plotted as a thin line.

# Results: Validation for 3/2 Power of Law

- Relationship between Non-dimensional wave height and period
- Windsea shows
  - good agreement with  $H^* \sim T^{*3/2}$
- Swell shows unclear relationship
  - scattered around  $H^* \sim T^{*2}$  ?



# Results: Wave Steepness and Directional Spreading

- Investigation about determinant for directional properties

$\overline{\sigma_\theta(f)}$ : frequency-mean directional spreading

$H_{m0}/L_p$ : wave steepness

## Windsea:

$\overline{\sigma_\theta(f)}$  increases with  $H_{m0}/L_p$ ,

directional spectrum becomes broad

as wave steepness increases

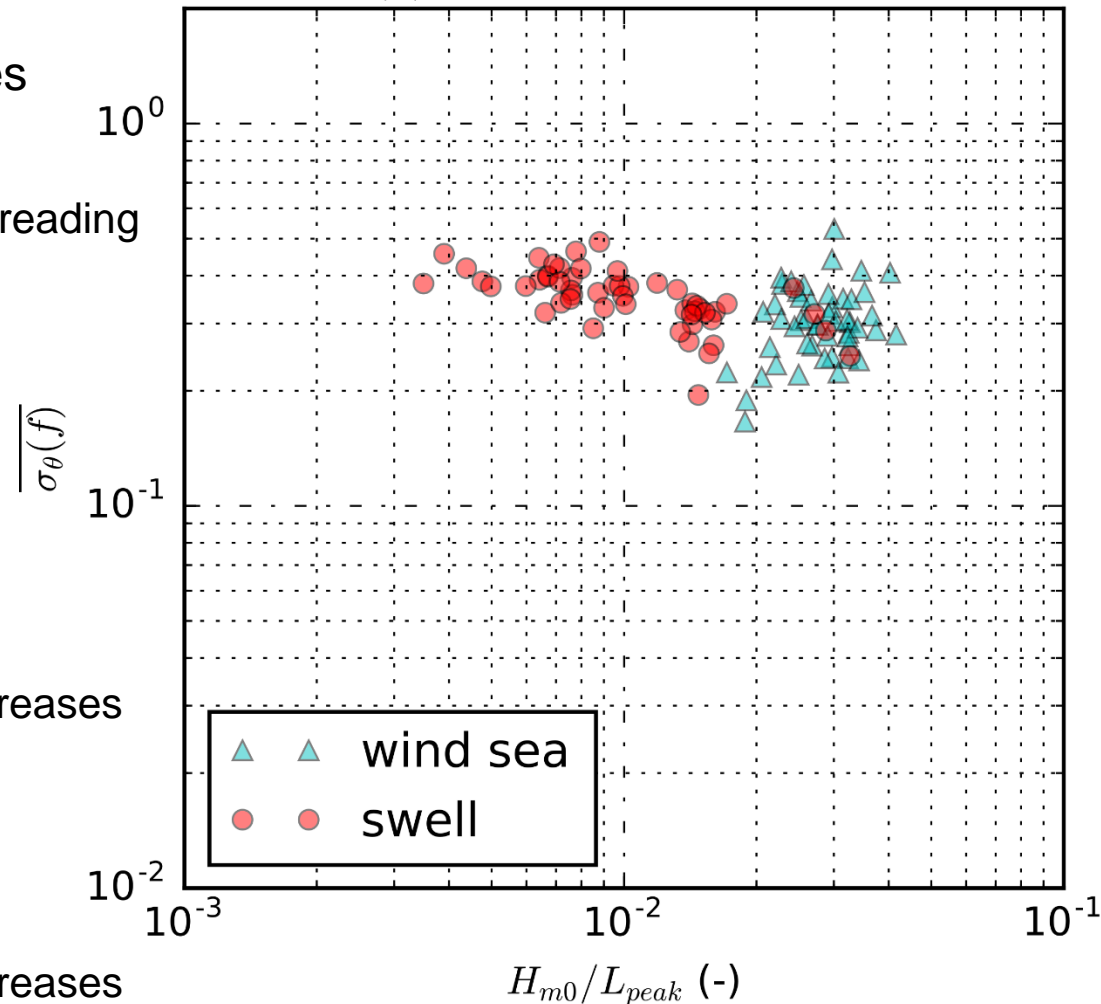
## Swell:

$\overline{\sigma_\theta(f)}$  decreases with  $H_{m0}/L_p$ ,

directional spectrum becomes narrow

as wave steepness increases

$\overline{\sigma_\theta(f)}$  vs wave steepness:



$\sigma_\theta(f)$ : circular root mean square spreading,  $\sigma_\theta(f) = \sqrt{2(1 - m_1(f))}$

$m_1(f) = \sqrt{a_1(f)^2 + b_1(f)^2}$ ,  $a_1(f) = \int \cos(\theta)G(\theta|f)d\theta$ ,  $b_1(f) = \int \sin(\theta)G(\theta|f)d\theta$

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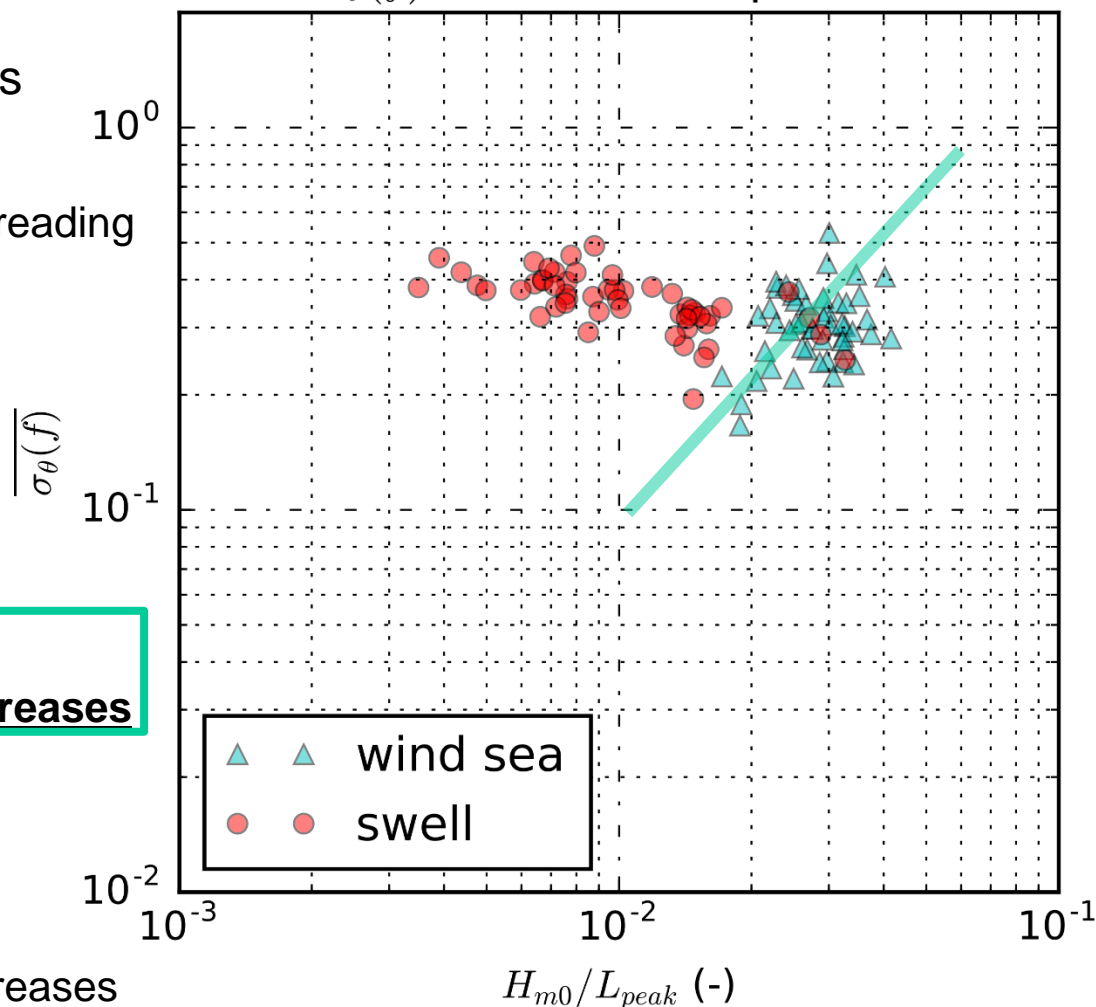
directional spectrum becomes **broad**  
as wave steepness **increases**

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$\overline{\sigma_\theta(f)}$  decreases with  $H_{m0}/L_p$ ,

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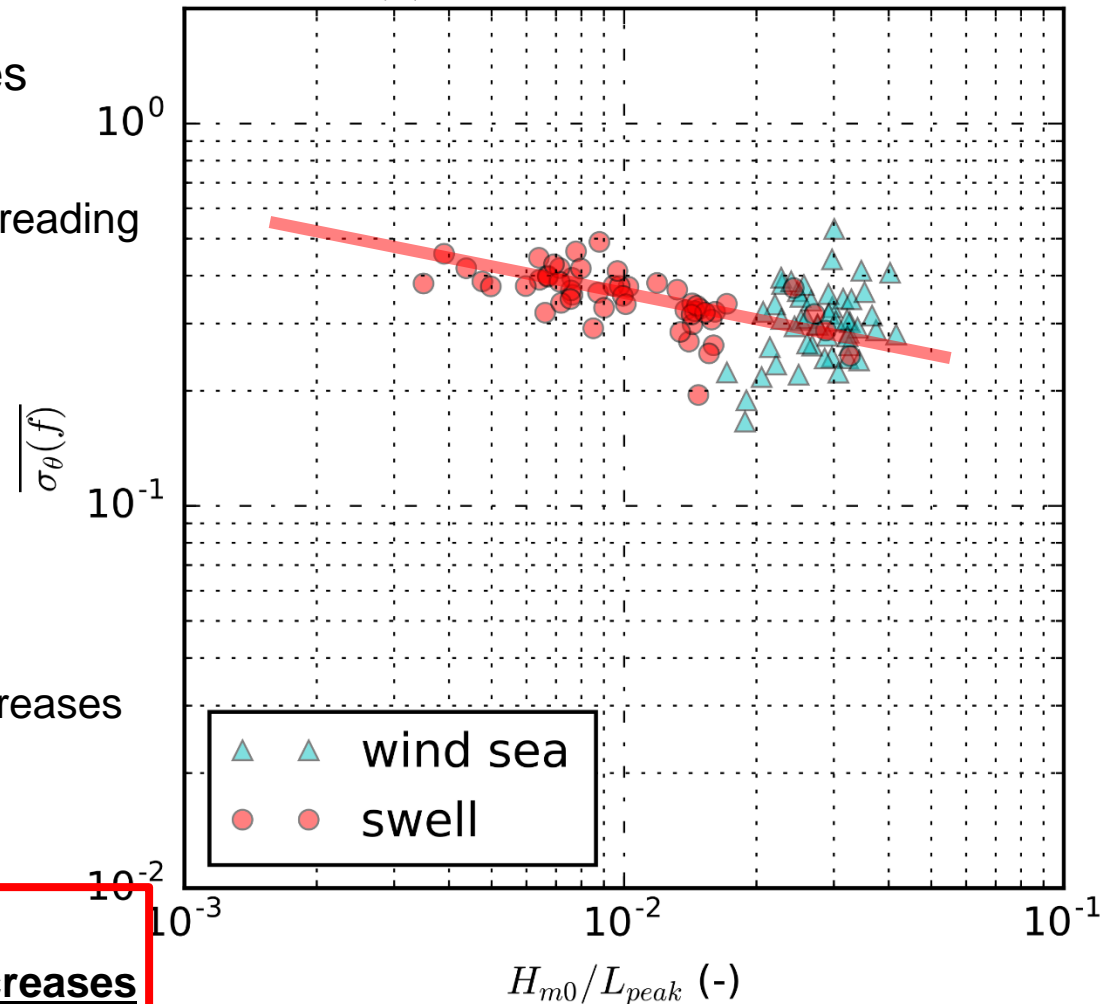
## Swell:

$\overline{\sigma_\theta(f)}$  decreases with  $H_{m0}/L_p$ ,

directional spectrum becomes **narrow**

as wave steepness **increases**

$\overline{\sigma_\theta(f)}$  vs wave steepness:



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# Results: Inverse Wave Age and Directional Spreading

- Investigation about determinant for directional properties

$\overline{\sigma_\theta(f)}$ : frequency-mean directional spreading

$U_{10}/C_p$ : Inverse wave age

## Windsea:

Concentrating in narrow range of  $U_{10}/C_p$

Difficult to interpret clear dependency

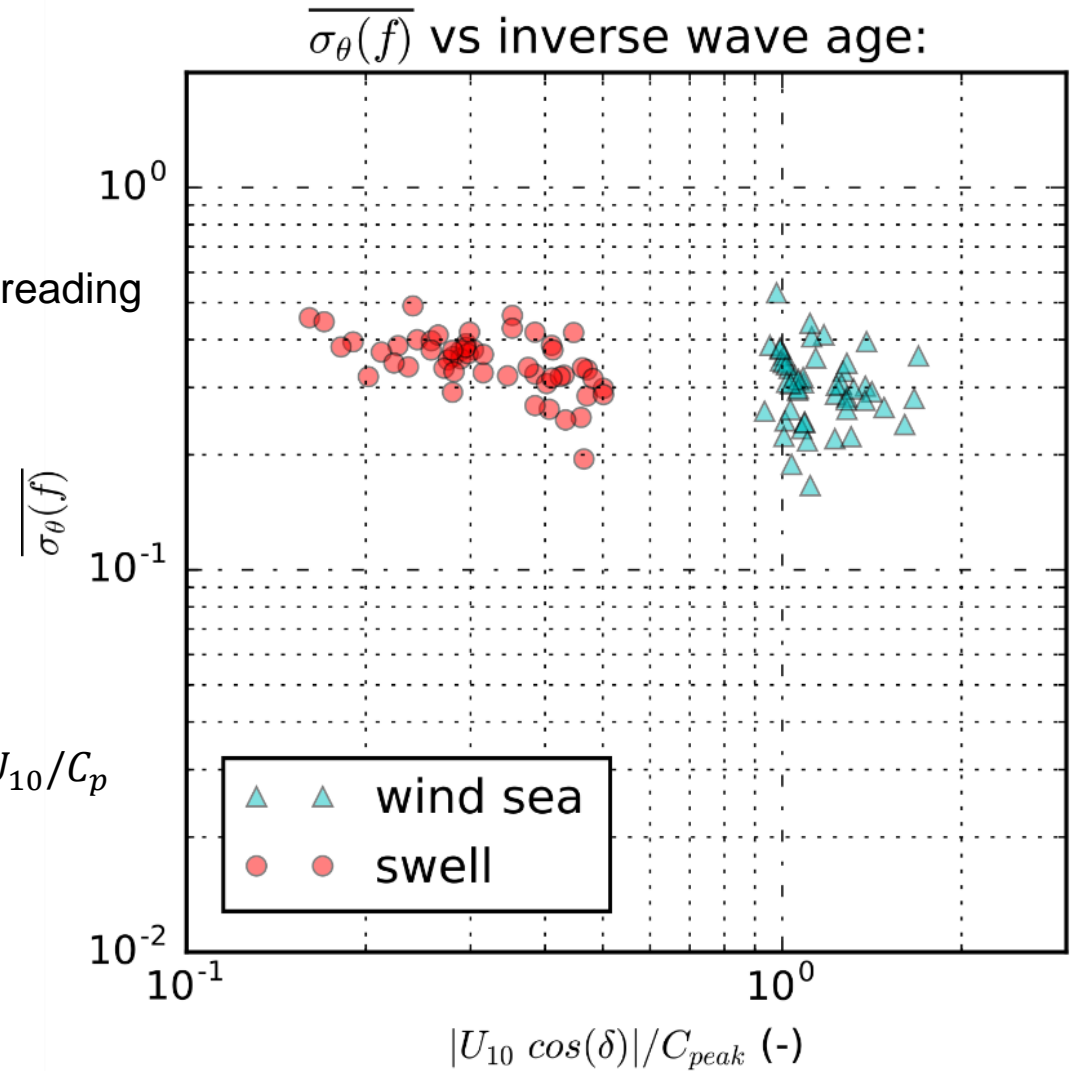
between  $\overline{\sigma_\theta(f)}$  and  $U_{10}/C_p$

## Swell:

$\overline{\sigma_\theta(f)}$  decreases with  $U_{10}/C_p$ ,

directional spectrum becomes narrow

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# Results: Inverse Wave Age and Directional Spreading

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$U_{10}/C_p$ : Inverse wave age

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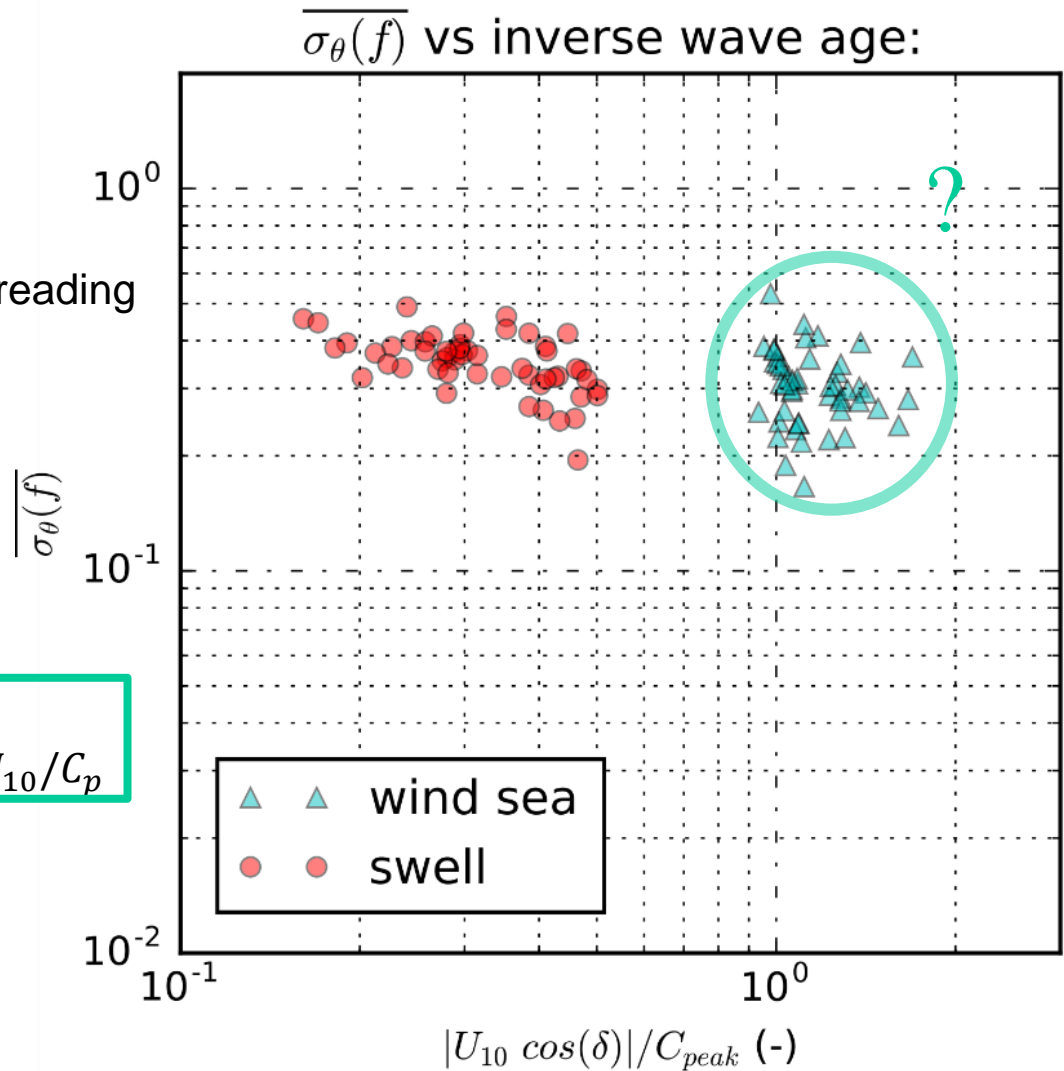
Difficult to interpret **clear dependency** between  $\overline{\sigma_\theta(f)}$  and  $U_{10}/C_p$

Swell:

$\overline{\sigma_\theta(f)}$  decreases with  $U_{10}/C_p$ ,

directional spectrum becomes narrow

as inverse wave age increases



# Results: Inverse Wave Age and Directional Spreading

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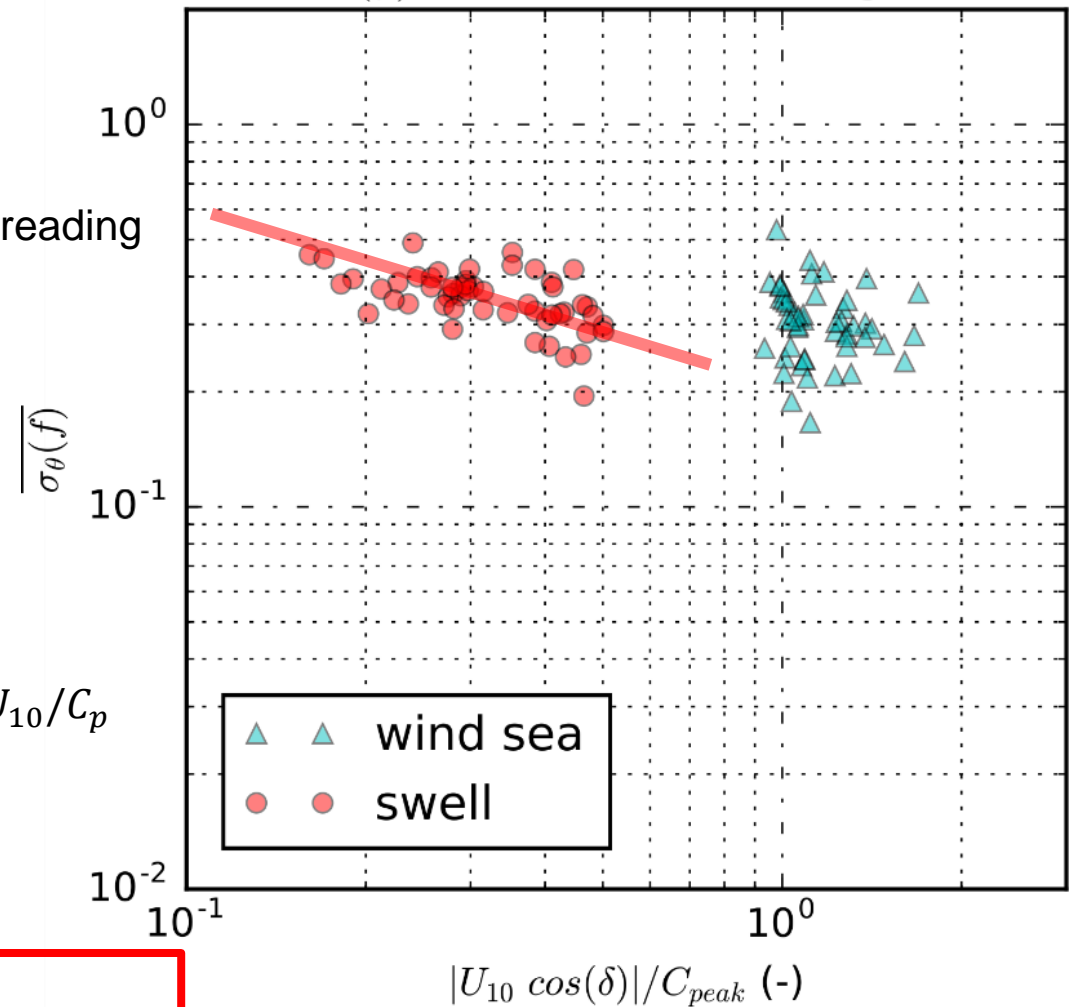
between  $\overline{\sigma_\theta(f)}$  and  $U_{10}/C_p$

Swell:

$\overline{\sigma_\theta(f)}$  decreases with  $U_{10}/C_p$ ,

directional spectrum becomes **narrow** as inverse wave age **increases**

$\overline{\sigma_\theta(f)}$  vs inverse wave age:



# Discussions and Conclusions

- Non-dimensional wave statistics is **limited in narrow range** in this field observations
  - Contrary to the previous study, ONLY the field data was investigated
- Wave steepness is **insufficient for separating windsea and swell**
  - Wave steepness of wind wave and swell can be overlapped
- Directional property of dependency on wave statistics is different in windsea and swell
  - windsea: directional spectrum becomes **broad** as wave steepness **increases**
    - Consistent with previous study, proposed by Goda et al. 1973
  - Swell: directional spectrum becomes **narrow** as wave steepness **increases**
    - Swell show the property contrary to windsea

# Future Works

- Validation of theories in the previous study with data in various scales
  - Focused in design use, research for extreme wave events, etc.
- Improvement on inversion method of directional spectrum
  - Result of inversion is NOT always perfect under too calm or severe sea state
- Improvement on data processing before inversion
  - In-situ data contains various type noise: spike, drift, jump, etc.
  - Careful processing including visual inspection is available for reliable analysis
  - Automated process to reject or modify such marginal data
    - is required for analyzing much data efficiently